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T A B L E des M A T I E R E S

1ère Partie

/ XVth GENERAL ASSEMBLY of the IUGG
 Moscow, 28 July - 14 August 1971

I - NEW ORGANISATION OF THE INTERNATIONAL ASSOCIATION OF GEODESY (I.A.G.)	p. 1-3
- Sections	
- Commissions	
- Special Study Groups	
- Duration of mandates	
- General Assembly	
II - DECISIONS CONCERNING THE SECTION III "GRAVIMETRY" (old Section IV)	p. 1-5
- Officers	
- Sp. Study Groups	
III - RESOLUTIONS and RECOMMENDATIONS	p. 1-6
- English text (Resolutions N°11-12-13-14-15-16 (Recommendations N°4-5-6-7)	
- Texte français (idem)	p. 1-10
IV - PROCEEDINGS of the MEETINGS of SECTION IV of the I.A.G. "GRAVIMETRY" by T. HONKASALO Secretary	p. 1-15
V - LIST OF SOME PAPERS read or distributed at the General Assembly, of interest for gravimetrists	p. 1-25
<u>ANNEXE A</u> - The INTERNATIONAL GRAVITY STANDARDIZATION NET 1971 (IGSN 71)	p. 1-29
Introduction (abstract)	
1-2 - Evolution of world gravity standards	p. 1-29
1-3 - Summary of the work of SSG 4.05	p. 1-32
<u>ANNEXE B</u> - Une TENDANCE de la VARIATION de la PESANTEUR observée au B.I.P.M Sèvres (France), par A. SAKUMA	p. 1-33
<u>ANNEXE C</u> - REPORT of the Sp. St. Gr N° 4.21 on "SPECIAL TECHNIQUES OF GRAVITY MEASUREMENTS" by T. HONKASALO (abstract)	p. 1-36

2ème Partie

TABLES for CONDENSATION REDUCTION
by J. C. BHATTACHARJI

p. T - 1

- TEXT

Introduction

Fundamental Tables for reducing condensations effects
at condensation surface of any elevation

p. T - 5

Special Tables for reducing condensation effects at
Earth Model and Geoidal surfaces

p. T - 7

- FUNDAMENTAL TABLES IN UNITS of 10^{-8} Gal

Zones A..... O_2
Zones 18.....1

p. T - 13

p. T - 14

p. T - 39

- SPECIAL TABLES

Table 1 - EFFECT E for EARTH'S TOPOGRAPHY $h-H$
condensed at EARTH MODEL
Compartment of zones A to 1

p. T - 45

p. T - 46

Table 2 - EFFECT E for EARTH'S TOPOGRAPHY $h-h_s$
condensed at GEOID
Compartment of zones A to O_2

p. T - 52

Table 3 - EFFECT E-T for EARTH'S TOPOGRAPHY $h-h_s$
with condensation at GEOID
Compartment of zones 18 to 1

p. T - 53

Table D - BULLARD'S CORRECTIONS

p. T - 54

Prof. P. TARDI
Dr. S. CORON

XVth GENERAL ASSEMBLY OF THE IUGG
(Moscow, 28 July - 14 August 1971)

- I -

NEW ORGANISATION OF THE INTERNATIONAL ASSOCIATION OF GEODESY (I.A.G.)

In the "Report of the General Secretary of the I.A.G." presented to the XVth General Assembly, M. LEVALLOIS give the results of the Cassiniis Commission created in Luzern (IUGG 1967), particularly to study the structures of the I.A.G. and namely to define the numbers and the content of the Sections, the outlines of their areas of interest, the junctions with the Commissions, the structure of the Study Groups...

We are extracting the principal points, which can be of special interest to the participants of the section "Gravimetry" :

Sections of I.A.G.

It was unanimously acknowledged that the division of the Association into sections was an excellent organisational set up and should be retained. After lengthy discussions, the following proposal for organisational chart was adopted, 5 sections are retained arranged as follows :

I - Nets

Triangulation : trilateration; bases, precision surveying, astronomical positioning; refraction; nets at sea.

II - Aerospace techniques

Geometric and dynamic aerospace techniques including optical measurements of distances and Doppler measurements, stellar triangulation ; application of interferometric techniques by means of long bases and lunar distances measuring techniques.

III - Gravimetry

Absolute and relative determination of the gravitational force ; gravity nets gravimetry at sea and in space ; measurements of the variations of the gravitational force : determination of the gravitational force gradients ; reduction, interpolation and extrapolation of the force of gravity..

IV - Theory and Data Processing

Mathematical geodesy ; application of potential theory and astromechanics ; applied statistical and numerical analysis ; adjustement theories ; methods of least squares ; data processing.

V - Physical Interpretation

Shape of the Earth, geometrical and physical parameters ; reference systems; variation in time including pole movement, rotation, earth tides and movements of the crust ; mean sea levels ; geological and geophysical consequences stemming from the geodetic results.

Commissions

The Cassinis commission approves the existence of commissions, they shall no longer be called "permanent", and, at least, as far as the technical Commissions are concerned, they should be under a Section.

The Cassinis Commission puts forward the following proposal :

"On a proposal by the Sections Chairmen, the Executive Committee appoints and disbards the Commissions. These decisions are approved by the General Assembly. The Chairman of a Commission is elected by the Council and the choice approved by the General Assembly".

Special Study Groups

The Cassinis Commission approves the existence of Special Study Groups. Each Group is under a Section and is formed for a specific study work, in principle lasting one term, (interval between two General Assemblies) and reports to its Section, during the General Assembly, on the work accomplished. It is then decided if the Group must carry on its work or be disbanded, there is no automatic reconstitution.

The Cassinis Commission, proposes that :

"On a proposal by the Sections Chairmen, the Executive Committee appoints and disbards the Study Groups and also nominates the Chairman".

Duration of mandates

The Cassinis Commission proposes that from now on, the duration of the Mandates of Sections Chairmen and Secretaries be for one term (the interval between two Assemblies) that is 4 years. It also proposes that in principle, Secretaries become Chairmen at the following election and this in order to ensure the continuity in the work of the Sections. It recommends that the Executive Committee be enlarged and made up of the Chairmen in Office and those of the previous term.

General Assembly

According to the new rules of the International Geodetic and Geophysical Union, Associations must during General Assemblies, hold joint meetings, common to at least two Associations. But furthermore, each Association will have the possibility of holding separate meetings during these General Assemblies and if it seems it necessary will hold a special General Assembly in between two General Assemblies of the Union. On this last point, the Cassinis Commission is unable to decide : some Geodesians staying away from Assemblies of the Union may be feared if special Assemblies are instigated.

DECISIONS CONCERNING THE SECTION III "GRAVIMETRY" (old Section IV)

Section "Gravimetry" remained almost unchanged, only the number IV was changed to III. The elections for the I.A.G. left also the section officers nearly unchanged.

- Officers in Section III, Gravimetry 1971-1975

President : Professor C. MORELLI, Osservatorio Geofisico Sperimentale Viale R. Gessi 4, 34123 Trieste, Italy

Secretary : Professor T. HONKASALO, Geodeettinen laitos Pasilankatu 43, 00240 Helsinki, Finland

- International Gravimetric Commission

President : Professor C. MORELLI (see above)

- Special Study Group 3.05

Establishment of a worldwide net of first order gravity stations and connections of these stations to the stations with absolute measured gravity.

President : Professor C. MORELLI (see above)

- Special Study Group 3.18

Absolute measurements of gravity

President : Professor A.H. COOK, F.R.S., Department of Geophysics, University of Edinburgh, 6 South Oswald Road, Edinburgh, Great Britain.

- Special Study Group 3.20

Measurements of gravity at sea

President : Professor J.L. WORZEL, Lamont Doherty Geological Observatory, Columbia University, Palisades, New York 10964, USA

- Special Study Group 3.21

Special techniques of gravity measurements

President : Professor T. HONKASALO (see above)

RESOLUTIONS AND RECOMMENDATIONS

We give here after the resolutions proposed by the Section "Gravimetry" and some other resolutions or recommendations concerning gravity but proposed by other Sections than IV.

Resolution n° 11

The IUGG,

recognizing that the Potsdam datum adopted in London in 1909, has served its purpose in providing a reference for international gravity measurements,

considering

- a) that for scientific purposes a more accurate system of gravity values is needed to provide both datum and scale,
- b) that the IAG has adopted at the Lucerne General Assembly in 1967 a provisional correction of - 14 mgal to the Potsdam value (Resolution n° 22),
- c) that the International Committee on Weights and Measures adopted in 1967 a resolution for a correction of - 14 mgal to the value of gravity in the Potsdam datum to be used for metrological purposes,
- d) that recent absolute, pendulum and gravimeter observations have provided a firm basis for the determination of datum and scale to the required accuracy,
- e) that the above mentioned measurements have been adjusted to provide a homogeneous International Gravity Standardization Net (IGSN 71) which defines the datum and scale and gives gravity values with the same order of accuracy throughout its range,
- f) that the establishment of such a system represents a major international effort, and provision for maintenance and improvement must be made,
- g) that the accuracy of the absolute determination of gravity is adequate for studies of variations in the distribution and displacements of masses, and variations of G,

recommends

- (i) that the International Gravity Standardization Net 1971 (IGSN 71) be adopted and published in the Bulletin Géodésique,
- (ii) that the Potsdam datum be corrected by the amount specified in the adjustment.



Among the remarks of Professor HONKASALO, we notice :
"quite confusing is, however, to give the name resolution for some of these and the name recommendation for some others, e.g. you have in the proof sheet resolution 12 and the recommendation 12. All must be resolutions and have final numbers ... I spoke with the Vice president of the IAG Professor KUKKAMAKI and he agreed with my opinion". (Letter Dec. 2, 1971).

Resolution n° 12

The I.A.G.,

recommends that a Permanent International Gravity Service, coordinated with the International Gravimetric Commission (IGC) and the International Gravimetric Bureau (IGB), be formed with the following functions :

1. to provide for expansion of the IGSN 71 to include areas where no stations now exist or values are not available,
2. the maintenance and re-observation of stations, and the replacement of any which may be destroyed,
3. to maintain the files in a computer-processable format and to incorporate new measurements into the existing files (one complete set of the values and descriptions to be deposited at the IGB),
4. to promote improvements in instrumentation, including the development of transportable absolute gravity meters, and their applicability to the standardization problem,
5. to establish in cooperation with the appropriate scientific institutions a net of permanent stations at which the absolute measurement of gravity, periodically repeated with an accuracy of the order of a few microgals, could be used as a geodetic reference and, in conjunction with other advanced geodetic methods, to monitor slowly varying parameters of the Earth,
6. to carry out computations necessary for incorporating new stations into the system,
7. to maintain contact with agencies active in the field of gravity measurements or using gravity data, to ensure that the IGSN 71 satisfies currents needs,
8. to provide advice, when requested, to agencies using the IGSN 71 in local standardization problems.

Resolution n° 13

The I.A.G.,

considering

- that an improved apparatus for the absolute measurement of g has been set up by the International Bureau of Weights and Measures,
- that the International Bureau apparatus offers the possibility to monitor the small variations of g arising from a variety of causes including secular effects of which tendency has been detected for the first time by means of absolute measurements,
- that this station (Sèvres, point A) has already been used as one starting point of gravimetric networks,

recommends

- that the installation at the International Bureau be maintained, and if possible improved, and that absolute measurements of g be performed as required in order to satisfy the needs mentioned above,

- that the International Bureau be requested to provide access to Sèvres not only to gravimeters but also other absolute gravity apparatuses to permit comparison of their results obtained at Sèvres point A against the value obtained with the International Bureau apparatus.

Resolution n° 14

The I.A.G.,

recognizing that Antarctica and the surrounding oceans are almost devoid of gravity observations and that gravimetric research is conducted by many groups from different nations,

recommends,

- a) the connection of all major Antarctic base stations to IGSN 71 as quickly as possible and with comparable precision,
- b) the establishment of a base net in the interior of Antarctica with stations located at appropriate intervals,
- c) the formation of a working group to plan and coordinate Antarctic gravity research and data exchange so that a precise and homogeneous coverage is obtained. Particular care should be given to the precision of geodetic coordinates associated with these stations.

Resolution n° 15

The I.A.G.,

recognizing that geodetic satellite and terrestrial data are of a great importance for studies of the Earth, its interior and various geophysical phenomena,

recommends that geodesists should actively participate in the Geodynamics project and similar programmes for investigating the structure and dynamics of the Earth's interior and their variation with time.

Resolution n° 16

The I.A.G.,

considering

- a) that it is necessary to improve the comparison between the elastic parameters of the earth found by measurements, and those of theoretical models,
- b) that the precision now obtained in the absolute measurement of g is in the order of one microgal,
- c) that experimental measurements of g , in order to be comparable, must be corrected for the effect of earth tides as well as for the indirect effect due to the ocean tides,

...

recommends

- a) that earth tides should be measured in those regions where measurements have not previously been made, with the aim of determining with precision the local constants, particularly in the southern hemisphere and eastern Asia,
- b) that continental profiles, already begun in the U.S.A. and Europe, should be extended as fully as possible in every region of the earth; these measurements must be carried out with instruments of the highest precision,
- c) to continue research on the disturbances due to the oceanic tides.

Recommendation n° 4

The I.A.G.,

recognizing the increasing importance of statistical methods for determining the terrestrial gravity field,

recommends research into the application of these methods to data of different kinds, and to their combination.

Recommendation n° 5

The I.A.G.,

recognizing the new problems raised by the processing of the contemporary highly accurate measurements of gravitational field elements by space and terrestrial techniques,

recommends that all relevant geophysical effects be investigated.

Recommendation n° 6

The I.A.G.,

recognizing the importance of astro-geodetic, gravimetric and other three-dimensional terrestrial methods for determining the figure and gravitational field of the Earth in detail and their variation with time, and

noting recent theoretical and practical work in the use of astro-geodetic and gravimetric methods for determining the geoid and geocentric positions,

recommends that the application of such methods and their combination with methods using satellite data be actively pursued.

Recommendation n° 7

The I.A.G.,

considering the great number of theoretical earth models currently used, and the accuracy obtained in experimental results during recent years,

recommends Section V of the I.A.G. to contact the I.A.S.P.E.I. in order to set up study group whose task should be to propose a single earth model to be used as a "Standard of comparison" for all results.

VOEUX (texte français)

Voeu n° 11

L'U.G.G.I.

reconnaissant que la valeur fondamentale de Potsdam adoptée à Londres (1909) a rempli son rôle en fournissant une valeur de référence pour les mesures internationales de pesanteur,

considérant

a) que les besoins scientifiques nécessitent un système plus précis de valeurs de la pesanteur qui fournisse à la fois la valeur et l'échelle,

b) que l'Association Internationale de Géodésie a adopté à l'Assemblée de Lucerne en 1967 une correction provisoire de - 14 mgal à la valeur de Potsdam (Voeu n° 22),

c) que le Comité International des Poids et Mesures a adopté en 1967 une résolution apportant une correction de - 14 mgal aux valeurs de la pesanteur dans le système de Potsdam, pour les usages métrologiques,

d) que des mesures absolues, au pendule et au gravimètre, ont fourni récemment une base solide pour la détermination des valeurs et de l'échelle avec l'exactitude nécessaire,

e) que les mesures mentionnées ci-dessus ont été compensées en vue de fournir un réseau gravimétrique international homogène d'étalonnage (appelé Réseau Gravimétrique International Unifié 1971) qui définisse les valeurs et l'échelle et qui donne des valeurs de la pesanteur avec le même degré d'exactitude dans toute son étendue,

f) que l'établissement d'un tel système est le résultat d'efforts internationaux considérables, et qu'on doit prendre des décisions en vue de son maintien et de son amélioration.

g) que l'exactitude de la détermination absolue de la pesanteur est adéquate pour l'étude des variations dans la répartition des masses et de leurs déplacements, ainsi que des variations de la constante de gravitation universelle G.

...

recommande

(i) que le Réseau Gravimétrique International Unifié 1971 soit adopté et publié dans le Bulletin Géodésique.

(ii) que la valeur de Potsdam soit corrigée de la quantité définie dans la compensation.

Voeu n° 12

L' A.I.G.

recommande que l'on constitue un Service International Permanent, travaillant en liaison avec la Commission Gravimétrique Internationale et le Bureau Gravimétrique International, chargé des tâches suivantes :

1. étendre le Réseau Gravimétrique International Unifié 1971 aux régions où il n'existe actuellement aucune station ou aucune valeur,

2. surveiller les stations, les moderniser et remplacer celles qui seraient détruites,

3. incorporer de nouvelles mesures dans les listes existantes et entretenir ces listes sous une forme compatible avec l'ordinateur (le Bureau Gravimétrique International centralisant valeurs et fiches),

4. provoquer le progrès de l'instrumentation et de son application aux problèmes d'étalonnage, en particulier par la mise au point de gravimètres absolus transportables.

5. établir, en coopération avec les institutions scientifiques adéquates, un réseau de stations permanentes dont les mesures absolues de pesanteur, répétées périodiquement avec une précision de quelques microgals, pourraient être utilisées comme référence géodésique et, associées à d'autres méthodes géodésiques récentes, pour surveiller les paramètres lentement variables de la Terre,

6. effectuer les calculs nécessaires pour l'incorporation de nouvelles stations dans le système,

7. rester en contact avec les Institutions travaillant dans le domaine des mesures de pesanteur, ou utilisant des données gravimétriques afin de garantir que le Réseau Gravimétrique International Unifié 1971 donne satisfaction aux besoins courants.

8. fournir des conseils, sur demande, aux Institutions utilisant le Réseau Gravimétrique International Unifié 1971 dans leurs problèmes de standardisation locale.

Voeu n° 13

L'A.I.G

considérant

- qu'un appareil perfectionné pour la mesure absolue de g a été construit par le Bureau International des Poids et Mesures.

- que cet appareil offre la possibilité de déceler les petites variations de g provenant de diverses causes, y compris les effets séculaires dont la tendance a été détectée pour la première fois par des mesures absolues,

...

- que cette station (Sèvres, point A) a déjà été utilisée comme point de départ de réseaux gravimétriques,

recommande

- que l'appareillage du Bureau International des Poids et Mesures soit maintenu et, si possible, amélioré et qu'on exécute les mesures absolues qui sont nécessaires pour satisfaire les besoins mentionnés ci-dessus,

- qu'on demande au Bureau International des Poids et Mesures de recevoir à Sèvres non seulement des gravimètres mais aussi des appareils de mesure absolue de la pesanteur afin que les résultats qu'ils fournissent à Sèvres, point A, puissent être comparés à la valeur obtenue avec l'appareil du Bureau International des Poids et Mesures.

Voeu n° 14

L'A.I.G.,

reconnaissant que l'Antarctique et les océans qui l'entourent sont presque dépourvus d'observations de pesanteur et que les recherches gravimétriques sont faites par plusieurs groupes de différentes nations,

recommande

a) qu'on relie les principales stations de référence de l'Antarctique au Réseau Gravimétrique International Unifié 1971 aussi rapidement que possible et avec une précision comparable à celle du Réseau International.

b) qu'on établisse un réseau de base à l'intérieur de l'Antarctique avec des stations convenablement espacées,

c) que l'on crée un "Groupe de Travail sur la pesanteur en Antarctique" pour établir les programmes de recherche et d'échange de données, et pour coordonner les travaux afin d'obtenir un réseau précis et homogène. Une attention particulière devra être donnée à la précision des coordonnées géodésiques des stations de ce réseau.

Voeu n° 15

L'A.I.G.,

reconnaissant l'importance des données géodésiques terrestres et de celles fournies par les satellites pour l'étude du globe, de l'intérieur de la Terre et de divers phénomènes géophysiques,

recommande que les géodésiens participent activement au projet géodynamique et à des programmes analogues comportant des recherches relatives à la structure et à la dynamique de l'intérieur de la Terre et à leur variation dans le temps.

Voeu n° 16

L'A.I.G.

considérant

a) qu'il est nécessaire d'améliorer les comparaisons entre les paramètres d'élasticité terrestres obtenus par des mesures et les modèles théoriques,

b) que la précision obtenue actuellement dans la mesure absolue de g est de l'ordre du microgal;

c) que les mesures expérimentales de g , pour être comparables, doivent être corrigées des effets des marées terrestres aussi bien que des effets indirects dus aux marées océaniques;

recommande

a) que des mesures de marées terrestres soient entreprises dans les régions où aucune mesure n'a encore été tentée, afin d'en connaître avec précision les constantes locales, particulièrement dans l'hémisphère Sud et la partie Est du continent asiatique;

b) que les profils continentaux actuellement en cours aux Etats-Unis d'Amérique et en Europe soient étendus aussi largement que possible à toutes les régions du globe terrestre et que ces mesures soient effectuées avec des instruments de la plus haute précision;

c) que les recherches sur les perturbations dues aux marées océaniques soient poursuivies.

Recommandation n° 4

L'A.I.G.,

reconnaissant l'importance croissante des méthodes statistiques pour la détermination du champ de la pesanteur terrestre,

recommande d'effectuer des recherches relatives à l'application de ces méthodes aux données de diverses natures et à leur combinaison.

Recommandation n° 5

L'A.I.G.,

tenant compte des problèmes nouveaux posés par le traitement des mesures de haute précision des éléments du champ de la pesanteur obtenues à la fois par des techniques spatiales et des techniques terrestres,

recommande l'étude de tout effet géophysique en relation avec ces mesures.

Recommandation n° 6

L'A.I.G.,

reconnaissant l'importance de la détermination de la figure de la Terre, de son champ détaillé de pesanteur et de ses variations dans le temps au moyen des méthodes astro-géodésiques, gravimétriques et d'autres méthodes terrestres tridimensionnelles, et

...

notant les récents travaux, théoriques et pratiques, relatifs à l'emploi de méthodes astro-géodésiques et gravimétriques pour déterminer le géoïde et des positions géocentriques,

recommande de poursuivre activement l'application de ces méthodes et de leur combinaison avec celles basées sur l'emploi des satellites.

Recommandation n° 7

L'A.I.G.,

considérant les nombreux modèles théoriques en usage actuellement et la précision des résultats expérimentaux de ces dernières années,

recommande à la section compétente de l'A.I.G de prendre contact avec l'A.I.S.P.I.T. afin de constituer une commission d'étude chargée de proposer un modèle terrestre unique qui serait utilisé comme "étalon de comparaison" de tous les résultats.

◦ ◦ ◦

PROCEEDINGS OF THE MEETINGS OF SECTION IV OF THE I.A.G. "GRAVIMETRY"

(Text of Professor HONKASALO completed with some abstracts).

Section IV held six meetings during the General Assembly at Moscow University. The chairman of all the meetings was Professor C. MORELLI and the secretary Professor T. HONKASALO. The agenda was as follows :

Meeting I. Tuesday, August 3rd at 11.30-13.00

1. Opening address by the president of Section IV, Professor MORELLI
2. Activité du Bureau Gravimétrique International, by Dr. S. CORON
3. Report of the SSG 4.05 "Establishment of a worldwide net of first order gravity stations and connection of these stations to the stations with absolute measured gravity", by the president of the SSG, Professor C. MORELLI.
Professor MORELLI gave a history of the work 1954-1971 ^{*} with particular reference to the last four years, during which the final adjustments to the International Gravity Standardization Net 1971 were performed. The account of the work was continued by the observers and members of the SSG.
4. Absolute gravity measurements in the IGSN 1971, by Professor J. FALLER.
Discussion on the possibilities of constructing a really portable absolute gravity meter weighing 100-150 kg (MORITZ, SAKUMA, FALLER, MORELLI).

Meeting II. Wednesday, August 4th at 15.00-16.30

Continuation of the SSG 4.05 report.

1. International cooperation, by Professor C. MORELLI
2. Pendulum observations and computings, by Professor T. HONKASALO.
3. Gravimeter observations, by Professor C. MORELLI
4. Adjustments and studies at the 1st Geodetic Survey Squadron, by C.T. WHALEN.
5. Computations at the Earth Physics Branch, by Dr. J. TANNER.
Discussion of the work of the SSG. Professor MORELLI drew special attention to the fact that the results of three independent adjustment computations, performed by C.T. WHALEN at the 1st Geodetic Survey Squadron, Professor U.UOTILA at Ohio State University and R.K. MCCONNELL and C. GANTAR at the Earth Physics Branch, differed only slightly.

* See : Introduction of the Report of the SSG 4.05, page 1 - 29.

Following discussion and studies of all systematic and individual differences in Ottawa, May 1971, the net now described is a new combined adjustment of 25 000 observations with grav., pend. apparatus and absolute devices. The new set of 1997 g-values obtained is no longer tied to a single reference point and is therefore assumed to possess almost the same degree of accuracy all over the world.

On the question of standard errors (TORGE), Mr. WHALEN referred to UOTILA's adjustment, which shows that all standard errors of gravity differences are smaller than ± 0.11 mGal.

On the question of value at Potsdam (ELSTNER) Professor MORELLI mentioned that the new value is more than 13.8 mGal smaller than the old one ; the new determination in Potsdam gives $\Delta g = -13.9 \pm 0.3$ mGal and agrees closely within the limits of its standard error .

Meeting III. Thursday, August 5th at 11.30-13.00

1. Further discussion about IGSN 71.

Professor BOULANGER proposed the publication of the work of SSG 4.05 in a complete form as soon as possible, and suggested that new independent accurate tying measurements be made between Potsdam and the most important gravity sites (England, Canada, USA, Australia) so that after careful new studies the proposal be submitted to the next General Assembly of the IAG.

The chairman postponed the decision to the next meeting of Section IV to have time to consider the arguments of both proposals, the adoption of IGSN 71 or postponement of it to the next General Assembly.

2. Une tendance de la variation de la pesanteur observée au Bureau International des Poids et Mesures, by Dr. A. SAKUMA *

Discussion (BOULANGER, FALLER, TERRIEN) on the accuracy of temperature and tidal corrections and the stability of the whole instrumentation.

3. Sur la nécessité d'un Service International de la Mesure absolue de la Pesanteur, by J. LEVALLOIS (x)

A permanent international service was proposed for studying the elements of the earth slowly varying with time (center of mass, seasonal variations of the radius, convection, variations of G).

The chairman supported the proposal but wanted to extend the scope of the permanent service, i.e. to take in the duties presented in the SSG 4.05 report. For absolute gravity studies he proposed a working group of experts (SAKUMA, FALLER, BELL, COOK).

* See : Annexe B, page I - 34

(x) See : Bul.Inf. n° 24, page I-42-43 "Quelques conséquences géophysiques des nouvelles méthodes de haute précision de mesures absolues de G."

4. Report of SSG 4.20, Measurements of gravity at sea, by M. TALWANI
Short report mostly of measurements and results in the Indian Ocean.
5. Vibration string sea gravity meter by Dr. C. BOWIN

The Woods Hole Oceanographic Institution has utilized the MIT vibrating string (VSA) sea gravity meter for the past four years. Experience gained and instrument improvements are summarized. An automatic data acquisition system incorporating a small digital computer (Hewlett-Packard 2114A) has been developed to record gravity, ship's velocity, magnetic bathymetric and date and time information. The W.H.O.I. system consists of the VSA gravity sensor mounted upon a Sperry Mark 19 Gyrocompass which in addition to acting as a stabilized platform provides velocity north and east components of ship's motion. The entire system is installed in a portable enclosure to enable easy transfer from one ship to another. The excellent quality of the data obtained has improved the ability to resolve the gravity field at sea and allowed the development of improved navigation processing schemes.

Meeting IV - Monday August 9th at ^{He} 15.00-16.30

- I. The chairman reported that and some members of the SSG had consulted with Professor BOULANGER about his proposal to postpone the adoption of the IGSN 71 and said that after checking the accuracy of the proposed net, Professor BOULANGER withdrew his proposal to postpone the decision.

The chairman put forward two resolutions in Section IV. These were unanimously adopted.

2. Report of SSG 4.21 "Special Techniques of Gravity Measurements" by Professor TH. HONKASALO *
3. Problems of measuring small gravity differences. by Professor E.GROTHEN :

The results were found in connection with experiments done mainly with LaCoste-Romberg gravimeters by recording gravity; problems specific for any instrument type will not be dealt with (like, e.g. temperature and pressure control, read out errors) because they can be overcome by specific devices. Calibration is a main problem. Nonlinearities are produced by (1) read-out and output system, (2) elastic system even though $f = dg/dz = \text{const.}$ for zero length spring where $z =$ output; when accuracy $< 10^{-3}$ is considered such effects are to be taken into account. Digitising, is used for avoiding low accuracy of recorders ($\pm 2.10^{-3}$). In the absence of

noise, precise calibration (inbuilt device for repeated calibration) is an alternative of linearity but filtering is, of course, problematic in case of time dependent (tidal) observations. Since beam deflections and corresponding diameter variations of spring are so small feedback meters seem to be necessary mainly because of hysteresis for time dependent measurements.

Vibrational perturbations (like microseisms, M) and order of magnitudes of corresponding CC-effects are discussed.

After-effects and hysteresis of nonfeedback, meters may not simply depend on gravity changes with time but seem to be affected by vibrations too. Since microseisms usually varies with seasons this fact could influence long time (tidal) observations.

Discussion of accuracy (HONKASALO), air pressure influence (BONATZ).

4. Results of high-precision gravity measurements, by. Dr. Cl. ELSTNER:

A special profile consisting of several gravity stations was chosen in order to detect an upper limit of the accuracy for the measurements with gravimeters and in view of the possible influence of slow variances in the distribution of masses inside the upper layers of the earth crust and in the surrounding of the observation points. For geological reasons the profile was erected in the middle part of the GDR. There an approximate homogeneous sedimentary situation is to be found. To avoid calibration errors the six observation points include a small gravity range of 16 mGal. The length of the line amounts about 250 km.

The construction of the points and first measurements of gravity are briefly described. An evaluation of their accuracy is given. Repeated gravity observations during the next years should give a first information on the amounts of the different disturbing effects in the microgal-range and on the detection of possible mass deviations in the ground. The work was performed in cooperation with other institutions in the German Democratic Republic.

5. Gravity variations in time, by. Dr. V.V. FEDYNSKY.

Translated paper of V.V. FEDYNSKY, A. Sh. FAITELSON, K.E. VESELOV, E.A. ASARKINA, B.A. KRASNOK.

Discussion of accuracy (TORGE). No single closure exceeds 0.1 mGal.

6. Investigations of secular gravity variations in Iceland, by Professor W. TORGE.

In 1938, 1965 and 1970 gravity measurements across the young volcanic zone of northern Iceland have been carried out along a west-east profile of about 100 km length, including also the zone of the western adjoining tertiary plateau basalts. The aim of the survey was to detect whether there exist any variations of gravity with time or not in the mobile zone of recent volcanism and open fissures.

About 30 stations have been established in 1938 by using one Thyssen-Schleusener gravity meter with accuracy of a few 0,1 mgal. These stations have been reconstructed and completed by about 60 stations in 1965, when the points were monumented too. The observations in 1965 and 1970 were performed with LaCoste and Romberg gravity meters, the standard deviation being less than 0.02 mgal.

The comparison of the gravity values obtained in the different periods indicates, that the gravity in the young volcanic zone might have more increased with time than in the tertiary basalts, the time rate over the total profile being about 0,01 mgal per year. Because of the strong correlation of the 1938 gravity values, the change between 1938 and 1965 is not significant, whereas between 1965 and 1970 a significant change is indicated.

Meeting V Tuesday, August 10 th at 16.30-18.00

1. Some problems of gravity measurements by means of gravimeters, by Dr. I. NAKAGAWA.
Report of the working group for gravity survey and precise levelling in Japan.

Formerly gravity measurements were independently made by the respective institution in Japan. From an urgent fundamental necessity of cooperative measurements of gravity with many gravimeters, a Working Group for Gravity Survey and Precise Levelling was organized under the Geodetic Society of Japan in 1969.

The Working Group had the symposium once a year and made precise gravity measurements over the areas of South Kanto, Tokai and Kii. An anomalous crustal activity has been observed in the areas of South Kanto and Tokai, and therefore simultaneous gravity measurements using 6 LaCoste gravimeters were carried out for investigating the second purpose described above. Simultaneous gravity measurements using 6 LaCoste and 2 Worden gravimeters were also carried out in the Kii Peninsula for investigating the first purpose.

A summary of conclusions obtained hitherto through the measurements and symposiums is as follow :

- (1) An extreme care must always be taken to the gravimeters.
- (2) Fundamental experiments about the gravimeters should systematically be made. Especially, effects of temperature, pressure and various shocks to the gravimeters should thoroughly be investigated for the gravimeters to be used.
- (3) Effect of mechanical, thermal and other shocks to the gravimeters is the most important in field measurements. Fluctuation amounting to 0.02 mgal was observable even by taking in and out the gravimeters to and from the carrying case.

- (4) Drift of the gravimeters is not linear even for a short duration of measuring time.
 - (5) It is rather difficult to keep a measuring accuracy always within 0.01 mgal even by careful measurements in quiet stations.
 - (6) An opportunity to detect the gravity change is expected for great earthquakes with magnitude larger than 7.5. For this purpose, gravity measurements should simultaneously be made with precise levelling.
 - (7) An improvement of the gravimeters is necessary in order to increase the measuring accuracy.
2. Pendulum apparatus with a constant amplitude pendulums, by Dr. V.A. ROMANJUK :
- (1) The most important feature of the apparatus is an electro-magnetic pusher with a reverse coupling for providing a constant amplitude for a long time. Experiments do not reveal the fact of the electro-magnetic pusher affecting the pendulum period within the accuracy the periods are measured with (1×10^{-9} sec).
 - (2) The pendulum elements are interconnected by means of springs. A brass glass rigidly fixed to the pendulum ensures temperature compensation.
 - (3) The pendulum swinging periods are registered photoelectrically. Physical periods of 2 actual and 1 fictitious pendulum are measured by the method of "typing" the pendulum pulses to quartz clock "minute" pulses.
 - (4) The amplitudes of 2 actual pendulums are equalized within $\pm 4''$. The amplitude is measured by means of the auto-collimation method with the error $\pm 0''.5$.
 - (5) Experimental forming of the cylindrical surface resting on the knife-edge of about 10 radius ensures the speed of the "drift" period equal to $3 + 8 \times 10^{-10}$ sec/hour compared to 50×10^{-10} sec/hour drift period of the non-formed knife edge.
 - (6) In 1969 the first experimental measurement of Δg 800 mgal with the error ± 0.02 mgal for actual pendulum was carried out.
- Discussion on accuracy and technical details
3. Measurements of the vertical gradient of gravity in mountains, by Dr. M. PICK :
- In the Czechoslovakian mountain test-area in the High Tatras, the investigation of the vertical gradient of gravity has been started. An area of 250 km² was covered by about 60 points, at which the vertical gradient of gravity was determined correctly from gravity measurements on two levels by means

of gravimeter Sharpe Canadian CG2. The procedure of measurements was arranged in such a way for the measurement errors to be less than 20 E. The influence of the gravity topographic corrections was reduced by a suitable procedure as well. At the same time the value of the vertical gradient of gravity was computed from the map of gravity anomalies. The accuracy of computed values dg/dz is lower than that of the measured values, and may amount to about 100 E.

4. Submarine vertical gradient of gravity, by Professor J. LAGRULA :

The vertical gradient of g in the oceanic masses, is a slowly decreasing function of the depth. A bathyscaphe allows to determine it. Its anomalies are more quiet than in continental areas and can therefore be used for geodetic purposes.

5. Reduction of vertical gradients of gravity, by Dr. H. MÄLZER :

The observed vertical gradients of gravity are dependent in a high degree on the topographic masses. Compared with the theoretical value differences of $\pm 30\%$ and more are existing.

Above all the topography of the immediate vicinity must be known for the reduction very exactly. Questions of the accuracies for heights and density and for the dimensions of the compartments are discussed. The influence of masses far from the observation point are investigated theoretically. The highest disturbances exist when the angle between the horizontal and the direction to the disturbing mass has $35^{\circ}3$. This result already found by N. KUMAGAI, E. ABE, and Y. YOSHIMURA (1962) is established in the computed table of reduction values.

The reduction of observed vertical gradients of gravity is illustrated by examples. The measurements are carried out with several LaCoste-Romberg gravimeters.

6. Experiment of the determination of the Cavendish gravity constant, mass and average density of the earth, by Dr. M.U. SAGITOV :

The exact expression for the momentum of the attractive forces of the trial masses is derived, and the unlinear theory of an experiment on the determination of the gravity constant is improved. Differents effects are considered, which should be taken into account in an exact experiment. It is described a laboratory installation for this experiment, which was completed recently. Preliminary results of determination of the Cavendish gravity constant, mass and average density of the Earth are presented and considered. The preliminary results promise an accuracy of 10^{-5} , theoretically 10^{-6} is possible.

Discussion : Why use cylinder masses, not balls ? Accurate cylinders are more easily prepared but the theory is more difficult

The chairman read the proposal of Professor BOULANGER and Dr. SAGITOV for a resolution :

" The International Association of Geodesy draws your attention to the necessity of improving the value of the constant of gravitation in the metric system of mass, length and time units as well as of the value of the Earth's mean density .

To distinguish this new constant of gravitation from other ones, such as the Gaussian, Einsteinian, geocentric and selenocentric, it is suggested to call it the Cavendishian constant of gravitation or merely the Cavendishian constant".

During the discussion (TERRIEN, FALLER, MORELLI) attention was drawn to the fact that the question of nominating a new constant in the metric system belongs to the International Committee of Weights and Measures and Several Scientific Unions, not only the IAG. Thus this resolution was not adopted by Section IV.

Meeting VI Wednesday, August 11th at 16.30-18.00

1. Some problems of tidal gravity correction, by Professor P. MELCHIOR.

A plan made by the Commission of Earth Tides to measure the tidal gravity constant with standardized gravimeters at several stations along two lines in Europe, one from Liverpool to Padova and the other from Lisbon to Helsinki was presented.

Discussion of local effects.

2. Les grandes anomalies de la carte gravimétrique de l'Afrique Occidentale et Centrale, by Doctor P. LOUIS.

(1) L'Office de la Recherche Scientifique et Technique Outre Mer a établi, durant ces quinze dernières années, une carte gravimétrique de reconnaissance qui couvre une grande partie de l'Ouest africain y compris le bassin du Tchad (Sénégal, presque toute la Mauritanie, le Mali, la Côte d'Ivoire, le Togo, le Dahomey, le Niger, la Haute Volta, le Tchad, le Cameroun et la République Centrafricaine). La densité des stations est d'environ une pour 70 km². Les résultats des levés ont été publiés à l'échelle de 1/1.000.000. Des cartes de synthèse (anomalies de Bouguer et anomalies isostatiques) ont été établies.

(2) La carte isostatique permet de constater que si globalement la compensation est réalisée à l'échelle de mille ou deux mille kilomètres, il n'en est plus de

...

même à l'échelle de quelques centaines de kilomètres. On constate, en effet, l'existence de grandes anomalies régionales. Ce sont les principales d'entre elles qui sont examinées et pour lesquelles des explications sont proposées dans cette communication...

3. About the preparation of a gravity map of the Western Section of Cuba at a 1:100 000, by G. OLIVA :

This paper describes, in some of its aspects, a recollection of gravity data concerning an area of approximately 40000 square kilometers in Western Section of Cuba, land and sea shelf being herein included. The paper has taken into account gravity surveys performed from 1929 to the present day, as well as the various recollections of data made during that period.

Sources used comprise documentation obtained from the Cuban Ministry of Mines, Fuel and Metallurgy, from the Cuban Institute of Geodesy and Cartography, and from a number of foreign companies, mainly American, that have performed geophysical surveys in Cuba before 1959.

Some 20000 gravity stations, taken from those diverse sources, have been checked for the area under study. That processing of the points selected guarantees compliance to all the conditions required in the preparation of gravity maps to a 1:100000 scale with a contour interval of 2 mgal. The methodology followed for the preparation of the map is here given in two variants ; Free-Air Anomalies and Bouguer Anomalies...

4. Airborne gravity, by O. WILLIAMS. :

Description of research in gravity measurements in the aeroplane and the helicopter.

5. Proposed International System Unit of Gravitational Acceleration.

The chairman reported on the content of Dr. W.I. REILLY's paper, in which he proposed a new name for unit m/s^2 = Galileo with the symbol G1.

In the discussion no reason was found for the new naming. Furthermore the symbol was found to be confusing, since it has already the meaning of Gigalitre (TERRIEN).

6. Gravity field, structure of the under-ice relief and the earth's crust in Antarctica, by Dr. N.P. GRUSHINSKY :

The survey of the bulk of gravimetric data on the Antarctic and that of the fundamental gravity connections are given. The general plan of the future study of gravity field of the Antarctica is suggested, which secures the maximal information for the geodetic and geophysical study of the continent.

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The fundamental connection between the gravity anomalies and the structures of the under-ice relief and Earth crust are discussed.

The conclusion on the high degree of the isostatic compensation of the Antarctica is done.

The discussion showed international cooperation to be necessary and a resolution on it was proposed. (Resolution n° 14).

Meeting of the SSG N° 4.21

Special techniques of gravity measurement.

Wednesday, August 11th at 9.30, Chairman Professor T. HONKASALO.

1. The most important task of this study group was stated to be high precision field measurements.
2. It was considered necessary for the SSG to have a large number of members to get information of high precision techniques all over the world. An active working group is necessary if joint test measurements are planned.
3. International cooperation was planned to remeasure the Fennoscandian land uplift line. This would provide an opportunity to test the influence of the improvements of new reading devices and thus separate the drift and reading uncertainties.

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SOME PAPERS READ OR DISTRIBUTED AT THE GENERAL ASSEMBLY
of interest for gravimetrists

- K.H. AFSHAR , H. ZOMORRODIAN : The measurements and the adjustments of the First Order Gravity Network in Iran.
- A.I.G. : Liste des stations de déviation de la verticale rattachées au Réseau Européen; liste n° 4, 1971.
- A.I.G. : Geodetic Reference System 1967, Bul. Géod.
- P. BANKWITZ, E. BANKWITZ : Conceptions about Motions along Oceanic Fracture Zones.
- G. BARTA : Interpretation of the Geoid Shape.
- G. BARTA : On the hypothesis of the secular variations of gravity field.
- G. BARTA : Über die Massenverteilung der Erde auf Grund der Geoidform.
- E. BAYER : Die Berücksichtigung des lunisolaren Effektes in der geodätischen Beobachtungen in Ungarn,
- Arne BJERHAMMAR : Prediction And Filtering of Non-Stationary Stochastic Processes in Geodesy.
- Arne BJERHAMMAR : Linear Unbiased Estimation Theory.
- M. BONATZ, P. MELCHIOR, B. DUCARME : Station : Longyearbyen (Spitsbergen) Mesures faites dans les trois composantes avec six pendules horizontaux VM et trois gravimètres Askania.
- L. BRAGARD : Potentiel de déformation superficielle produit par un bourrelet de marée. Relations entre les nombres de Love h_n et k_n .
- R. BREIN : An instrument for precise measurements of small differences of the gravity.
Report of the Institut für Angewandte Geodäsie. Abteilung II des Deutschen Geodätischen Forschungsinstituts for the XVth General Assembly of the IUGG - IAG 1-14th August 1971 in Moscow.
- Milan BURSA : Fundamental Geodetic Parameters of the Earth's Figure and the Structure of the Earth's Gravity Field Derived from Satellite Data.
- Milan BURSA : Determination of parameters of a selenocentric Reference System and Deflections of the vertical at the Lunar Surface.
- J. BYL : Results of tilt observations at Potsdam.

- T. CHOJNICKI : Problèmes des moyens d'élaboration des observations de marée en tenant compte des techniques contemporaines d'observation et de calcul.
- B.H. CHOVITZ : Generalized Three-Dimensional Transformations.
- G. DESVIGNES : Geoïde sur l'Europe et l'Atlantique Nord-Est.
- H.M. DUFOUR : La représentation du potentiel terrestre par des fonctions régionalisées.
- Cl. ELSTNER : Results of high precision gravity measurements.
- V.V. FEDYNSKY, A.Sh. FAITELSON, K.E. VESELOV, E.A. ASARKINA, B.A. KRASNOV : Gravity variations in time.
- F. FELDBACHER : Anwendungen einer numerischen Integrationsmethode bei der Lösung der Hauptaufgabe über grosse Entfernung.
- I. FISCHER : Recent Geoid Studies (Secretary's Report, Section V, 1967-1971).
- D. FLACH, O. ROSENBACH, H. WILHELM : Comparative harmonic analysis of records of 2 Askania borehole tiltmeters.
- D. FLACH, G. JENTZSCH, O. ROSENBACH, H. WILHELM : Ball-calibration of Askania borehole tiltmeters.
- J. FLICK, P. MELCHIOR, J-M van GILS : Le laboratoire souterrain de Geodynamique de Walferdange-Luxembourg.
- E.M. GAPOSCHKIN ym. : Geodetic Studies at the Smithsonian Astrophysical Observatory
- E. GROten : New determination of the gravitational constant.
- E. GROten : Problems in measuring small gravity differences.
- Rolland L. HARDY : Geodetic Applications of Multiquadric Analysis.
- T. HONKASALO : Report on IAG Special Study Group No.4.21 of Special Techniques of Gravity Measurements.
- T. HONKASALO : Geophysical Interpretation of Gravity Anomalies 1963-1970.
- Vladimir Hristov : Proposal by the Bulgarian National Committee on Geodesy and Geophysics for a Specification of the Geodetic Reference System 1967.
- Vladimir Hristov : Paramètres initiaux antérieurs et actuels en géodésie.
- K. HUBENY : Die exzentrische Anomalie als Parameter des Rotationsellipsoïds.
- Lubomir KUBACEK : Some mathematical aspects in the designing of geodetic nets.
- J. LAGRULA : Equal-area blocks.

- J. LAGRULA : Sur les corrections topographiques, géologiques et isostatiques.
C.R. Acad. Sc. Paris, t.272, p. 816-820 , 1971.
- J. LAGRULA, S. CORON : Submarine vertical gradient of the gravity.
- J.J. LEVALLOIS : Sur la nécessité d'un Service International de la Mesure absolue de la Pesanteur.
- J.J. LEVALLOIS : Calcul du géoïde gravimétrique sur le territoire de la France.
- J.J. LEVALLOIS : Sur le prolongement analytique du développement en harmoniques sphériques du potentiel terrestre.
- J.J. LEVALLOIS : Le système de Référence 1967.
- M. MAJEWSKA : The calibration of sharp gravimeter No. 228 G by means of tilting method in relation to temperature, atmospheric pressure and time.
- H. MALZER : Reduction of vertical gradients of gravity.
- R.S. MATHER, B.C. BARLOW, J.G. FRYER : A Study of the Earth's Gravitational Field in the Australian Region.
- R.S. MATHER : Practical Techniques for the Establishment of a World Geodetic System from Gravity Data.
- P. MELCHIOR : Interrelation between Ocean and Earth Tides.
- P. MELCHIOR : Rapport sur les Marées Terrestres 1967-1971.
- V. MITROVIC, M. GRASIC : Gravimetric calibration measurements in 1967 and 1969 (Yugoslavia).
- C. MORELLI, C. GANTAR, TN HONKASALO, R.K. McCONNELL, B. SZABO, J.G. TANNER, U. UOTILA, C.T. WHALEN : The International Gravity Standardization Net 1971 (IGSN 71).
- Helmut MORITZ : Series Solutions of Molodensky's Problem (Dtsche Geod. Komm. A:70).
- Foster MORRISON : The Manipulation of Density Layer Models for the Geopotential.
- Abdulah MUNINAGIC : Investigation of Real Geoid in Yugoslavia.
- I. NAKAGAWA : Some remarks on gravity measurements by means of gravimeters.
- G. OLIVIA : About the preparation of a gravity map of the Western Section of Cuba at a 1:100 000 scale.
- O.M. OSTACH : Solution of Stokes' Problem for a Boundary Ellipsoidal Surface by Means of Green Functions.
- L.P. PELLINEN, O.M. OSTACH, E.M. ORLOVA : Some Results of Astro-Gravimetric Leveling in the USSR.
- M. PICK : On the computation of the gravity terrain correction in Czechoslovakia.

- M. PICK : Measurements of the vertical gradient of gravity in mountains.
- M. PICK, J. PICHA : The Mean Density of the Earth.
- O.J. RAICES VIDAL : A Unified Procedure for the Solution of the Least Squares Problem and for Estimating the Precision of the Adjusted Variables.
- R.H. RAPP : Preliminary Report of the Equal Area Block Working Group.
- W.I. REILLY : Proposed international system unit of gravitational acceleration.
Geophys. J.R. astr.Soc. (1971) 22
- K. RINNER : Berich über die Erdezeitenstation im Grazer Schlossberg.
- A. SAKUMA : Une tendance de la variation de la Pesanteur observée au Bureau International des Poids et Mesures, Sèvres, France.
- A.Z. SAZONOV : Astro-Geodetic Network of the USSR. Lay-out and Adjustment Principales, Actual Accuracy.
- M. SCHADLICH : The Accuracy of Schematic Extended Astrogeodetic Networks.
- M. SCHADLICH : Modeltheory and the Technique of Geodetic Measurements - A General Conception of the Theory of Measurements.
- A. SCHLEUSENER, W. TORGE : Investigations of secular gravity variations in Iceland.
- M.M. SCHNEIDER : Measurements of the Earth's Tides at Vostok station in Central Antarctica.
- R. SCHULER, G. HARNISCH, H. FISCHER, R. FREY : Absolute Schweremessungen mit Reversionspendeln in Potsdam 1968-1969. Veröff. d. Zentralinstituts Physik der Erde No.10.
- D. SIMON : First results of earth tidal observations in airtight measuring boxes.
- Z. SIMON, L. TRÄGER : Polygons for studying the secular variations of the acceleration of gravity on the territory of Czechoslovakia.
- J.G. TANNER, R.A. Gibb: Gravity measurements in Canada January 1, 1967 to December 31, 1970.
- P. TARDI, S. CORON : Activité du Bureau Gravimétrique International, Paris.
- TEHRAN UNIVERSITY : Report on the Tidal Measurements of the Crust and Their Harmonic Analysis in Tehran.
- Erik TENGSTROM : Studies of different methods of Physical Geodesy for determining the figure of the Earth, and its external gravitational potential.
- Urho A. UOTILA : Interpolation and Extrapolation of Surface Gravity Values.
- A.P. VENEDIKOV, P. VARGA : Facteurs moyens des marées terrestres.

- H.W. WIRTH : Isostasy and Deformations of the Equipotential Surfaces in the Earth's Interior.
 - G. WITTLINGER, R. LECOLAZET : Sur les observations de marée clinométrique dans un long tunnel.
 - V.F. YEREMEEV, M.I. YOURKINA : L'expression des altitudes du Quasi-Geoïde et des déviations de la verticale pour le champ de Référence de l'Ellipsoïde de niveau.
 - H. ZOMORRODIAN : The measurements and the adjustments of the second order gravity network in Iran.
 - H. ZOMORRODIAN : A summarizing report on the establishment of the Iran National Gravity Calibration Line (I.N.G.C.L.) and the computation of the scale factors of the 4 applied LaCoste-Romberg gravimeter.
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ANNEXE A

THE INTERNATIONAL GRAVITY STANDARDIZATION NET 1971 (IGSN 71)

INTRODUCTION

..... 1.2. Evolution of world gravity standards

Early development

The first internationally accepted gravity reference system was known as the Vienna Gravity System. It was adopted in 1900 at the XIIIth Conference of the International Association of Geodesy held in Paris and had an estimated relative accuracy of ± 10 mGal.

The Potsdam Gravity System was introduced soon after (Borass, 1911) and internationally accepted at the 1909 meeting of the IAG in London. The relative accuracy of this system was estimated at ± 3 mGal and it corrected the Vienna System by - 16 mGal. Absolute gravity measurements in the past few decades indicated an error of + 12 to + 16 mGal in the absolute Potsdam value (Dryden, 1942; Morelli, 1946b; Berroth, 1949; Preston Thomas *et al.*, 1960; Cook, 1965a).

By the end of World War II it was evident that not only more absolute measurements but also an interconnecting net of relative gravity ties were required to define a new gravity system.

The publication of two independent adjustments (Morelli, 1946a; Hirvonen, 1948) demonstrated the insufficient distribution and accuracy of the existing gravity measurements and induced G.P. Woollard to undertake his pioneer work of promoting the geodetic use of new instruments (Worden and LaCoste Romberg gravimeters, Gulf pendulums) to provide new and more accurate relative gravity measurements all over the world (Woollard and Rose, 1963).

By the early 1950's many agencies had become involved in long range gravimeter and pendulum measurements. Inhomogeneity in the distribution of stations, in observational criteria and techniques and in data reduction methods necessitated the coordination of these activities. Accordingly action was taken through the IGC of the IAG and in 1954 Special Study Group 4.05 was formed with responsibilities in the following fields.

- (a) Absolute measurements of gravity
- (b) Network connections
- (c) International Gravity Formula.

A network of 34 stations to be known as the First Order World Gravity Net, were chosen at the 1956 meeting of the IGC in Paris and plans were made to concentrate on gravity connections between them. These stations were :

Algiers	Khartoum	Oslo
Azores	Kyoto	Ottawa
Bad Harzburg	Leopoldville	Panama
Beirut	Lisbon	Paris
Buenos Aires	Madison	Potsdam
Capetown	Madrid	Quito
Christchurch	M'Bour-Dakar	Reykjavik
Fairbanks	Melbourne	Rio de Janeiro
Helsinki	Mexico City	Singapore
Honolulu	Milan	Teddington
Johannesburg	New-Delhi	Vancouver
		Washington

During the next 8 or 9 years, relative gravity measurements were carried out on a world-wide basis by many observers. For logistic and technical reasons it was often impractical to carry out direct interconnection of the FOWGN stations and a large number of inhomogeneous subnets came into existence. Apart from partial solutions such as the basic world-wide work of Woppard and Rose (1963) and the European Calibration System (Kneissl, Marzahn, 1963b) the evolution of a new world gravity system proceeded slowly. Some significant improvements in instrumentation occurred during this period, e.g. development of the LCR gravimeter, but some disturbing aspects of the performance of pendulum apparatuses came to light. The state of the art was summarized (Morelli, 1963a) as follows :

- " a) Modern pendulum results, previously considered fairly reliable, have revealed (Woppard and Rose, 1963) tares and creep: that is, the same weak points as the gravity-meters; pendulums should therefore always be used in connection with properly chosen and studied gravity-meters;
- b) Results with modern, geodetic gravity-meters have shown that it may also be possible to measure large differences in gravity, provided that :
 - (1) they are operated in groups to control tares and creep;
 - (2) they are properly checked in the laboratory and on the calibration lines to detect and evaluate pseudo-periodic errors and nonlinearity;
 - (3) They are calibrated (on sufficiently accurate and extended calibration lines) to determine their scale-factor function.

It is well known that normally (3) can be done only by comparison with pendulum measurements.

It would seem that a solution would be impossible : we need the gravity-meters to be sure of the pendulums, and the pendulums to calibrate the gravity-meters.

Without proper consideration of this point, it is clear that confusion and uncertainty have been and will continue to be created. This is the reason that we do not yet have a "World Standardization System", although we have many "Calibration Systems" : referred to different pendulum apparatus, or to different pendulum results, or to a different evaluation of the same pendulum measurements".

The above situation led the SSG 4.05 during the XIII IUGG General Assembly (Berkeley, 1963) to propose, and the IAG to accept, a new philosophy as follow :

I. The establishment of three International Calibration Lines, with large gravity intervals.

They were : the ACL from Ushuaia to Point Barrow;
the EACL from Capetown to Hammerfest;
the WPCL from Christchurch to Sapporo.

II. The selection of a few stations on each International Calibration Line to be occupied by the best available pendulum apparatuses using the same operational criteria.

The chosen apparatuses were :

The Gulf pendulums;
The Cambridge " ;
The CGI " ;
The GSI " .

III. Groups of LaCoste and Romberg gravimeters would be observed on the Lines, for checking and calibration purposes.

IV. Interconnection would be established between the Lines for strengthening the structure and establishing the skeleton of the world net.

In the next few years, most of the measurements required for a new reference system were carried out. Important contributions were made by 1381st GSS with the first systematic global LaCoste and Romberg gravimeter measurements and by AFCRL who supported many projects, especially pendulum and absolute measurements.

The further coordinate the collection and reduction of data for the preparation of a final adjustment, the SSG 4.05 formed a sub-group of specialists devoted to the solution of specific problems. The representatives of the institutions most actively participating met for the first time in 1965 in Torino, Italy, and formed a permanent Working Sub-Group.

In the next few years the work of collecting and editing the new observations was carried out and many partial adjustments performed by members of the Working Sub-Group or others (TORGE, 1966). The increased relative accuracy of the best pendulum apparatuses (± 0.2 mGal and gravimeters (± 0.05 mGal) made possible the attainment of a world network of high relative accuracy. The problem of obtaining high absolute accuracy as well was solved with the development of Cook's semi-transportable apparatus (± 0.1 mGal) *

The amount of new data and the dangers inherent in continuous disseminations of new values led the International Gravity Commission at its last meeting (Paris, 1970) to request that only one final adjustment should be published and internationally adopted. The present report describes the adjustments and analyses which led to the final adjustment and the resulting International Gravity Standardization Net 1971.

This proposed system is estimated to have an absolute accuracy of ± 0.2 mGal or better over the gravity range of the earth.

1.3 Summary of the work of SSG 4.05

Reports of the SSG 4.05 have been presented and discussed at the following meetings :

1954, in Rome	- X IUGG General Assembly	(unp.)
1956, in Paris	- IGC Meeting	(unp.)
1957, in Toronto	- XI IUGG General Assembly	(see Morelli, 1959)
1959, in Paris	- IGC Meeting	(unp.)
1960, in Helsinki	- XII IUGG General Assembly	(unp.)
1962, in Paris	- IGC Meeting	(unp.)
1963, in Berkeley	- XIII IUGG General Assembly	(unp.)
1965, in Paris	- IGC Meeting	(unp.)
1967, in Lucerne	- XIV IUGG General Assembly	(unp.)
1970, in Paris	- IGC Meeting	(unp.)

Meetings of the Working Sub-Group have been held in :

- Torino,	April 1965;
- Paris,	September 1965;
- Bedford,	April 1967;
- Paris,	September 1970;
- Ottawa,	Mai 1971.

Some information concerning the Group's activity and discussions at the IGC and IUGG Meetings are printed in the "Bulletin d'Information" of the IGB.

The members of the Sub-Group participating in the final adjustment and the presentation of this report are :

AFCRL	: B. SZABO	1st GSS	: C.T. WHALEN
EPB	: R.K. McCONNELL, J.G. TANNER	OGST	: C. GANTAR, C. MORELLI
GL	: T HONKASALO	OSU	: U. UOTILA

* Faller's transportable apparatus (± 0.05 mGal) and Sakuma's stationary apparatus (± 0.03 mGal-systematic errors included).

Professor WOOLLARD contributed to the efforts of this Working Sub-Group, but could not participate in the final adjustment or in the presentation of this report. Professor WOLF acted from time to time as a consultant to the Sub-Group.

The IGSN 71 has been achieved within the framework of SSG 4.05 and through the cooperation of a great many agencies throughout the world.

ANNEXE B

UNE TENDANCE DE LA VARIATION DE LA PESANTEUR OBSERVEE
AU BUREAU INTERNATIONAL DES POIDS ET MESURES, SEVRES, FRANCE
PAR A. SAKUMA

Depuis Août 1967, les mesures de g ont été poursuivies périodiquement au B.I.P.M. par un gravimètre absolu de haute sensibilité (1) (2) et nous avons déjà signalé (3) que les valeurs moyennes de g observées pendant les trois premières années (Août 1967 - Août 1970) restaient constantes à environ $\pm 1 \times 10^{-8}$ ($\pm 10 \mu\text{Gal}$) près. Avec l'augmentation de la précision de nos mesures, les variations observées, qui atteignent par exemple 2×10^{-8} entre Août 1970 et Juillet 1971 doivent être considérées comme significatives. Nous ignorons les causes de ces variations et malgré nos efforts, aucun défaut instrumental de cet ordre n'a été décelé jusqu'à présent dans notre appareil.

(1) Bull. Géod. N° 69, 249/260, 1963.

(2) Bull. Information de l'I.G.N., Paris, N° 11, 8/19, 1970.

(3) Bull. Géod.. N° 100, 159/163, 1971.

Les valeurs moyennes de g au point de mesure A2 après correction théorique de l'effet luni-solaire (établie grâce à la collaboration du Professeur LECOLAZET de Strasbourg) et les écarts types d'une mesure sont pour la quatrième année :

	(980 925 657,4 \pm 2,0)	μGal	(Août	1970)
	(45,3 \pm 2,2)	"	(Octobre 1970)
	(86,3 \pm 9,4)	"	(fin Décembre 1970)
	(79,6 \pm 6,0)	"	(Janvier 1971)
$g_{A2} =$	(74,7 \pm 9,7)	"	(Février 1971)
	(71,0 \pm 8,7)	"	(Mai 1971)
	(80,1 \pm 6,2)	"	(Juin 1971)
	(76,7 \pm 5,5)*	"	(Juillet 1971)

La fig. 1 présente également tous les résultats de même nature obtenus depuis quatre ans.

Ce prototype de gravimètre absolu est composé actuellement par une centaine d'appareils. Les manipulations et l'entretien de ces appareils sont assez complexes. Néanmoins, les améliorations apportées jusqu'à présent permettent de mesurer g par un seul opérateur, après une heure de préparation, à la cadence d'une mesure de g toutes les 10 à 15 minutes, sauf pendant les tremblements de terre.

Ce gravimètre pourra être utilisé, dans l'avenir, pour plusieurs études géophysiques; vérification de la marée gravimétrique théorique, l'effet de l'affaiblissement du sol sous la surcharge atmosphérique, corrélations entre les dérives de g et les mouvements du pôle, etc.

* Résultat provisoire.

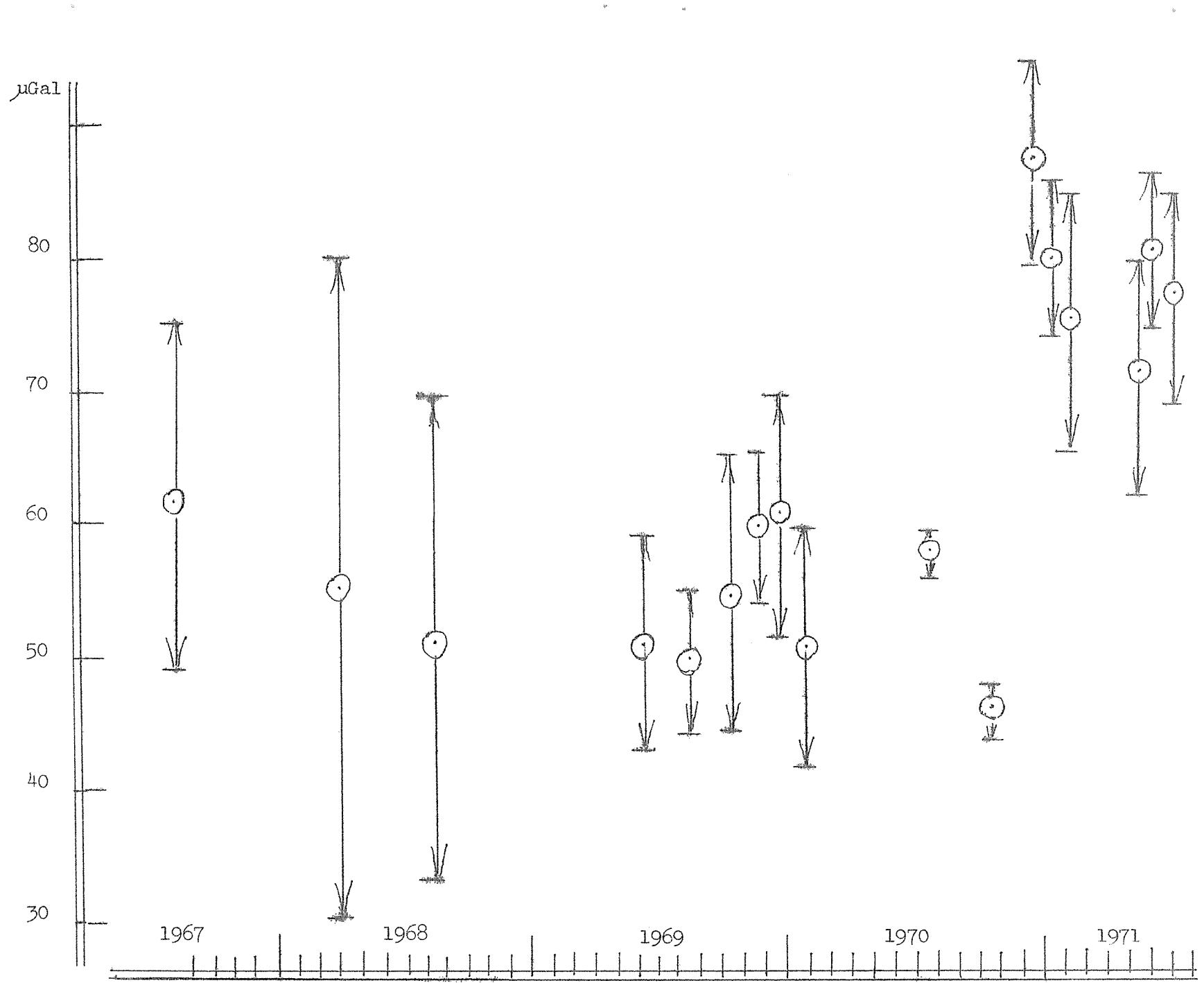


Fig. 1 - Valeurs moyennes de g avec les écarts types d'une mesure à
SEVRES Point A2 ; $g_{A2} = 980\ 925\ 600\ \mu\text{Gal} + \text{les valeurs en ordonnées}$

ANNEXE C

REPORT OF IAG SPECIAL STUDY GROUP N° 4.21
ON "SPECIAL TECHNIQUES OF GRAVITY MEASUREMENTS"
T. HONKASALO

A provisional Report was presented to the IGC (Paris, September 1970) and published in Bulletin Information N° 24 (I-52).

As the Report presented to the General Assembly in Moscow (August 1971) was only partly changed, we give hereafter some complementary information extracted of the last Report.

1 - INTRODUCTION

To study the secular variation of gravity, to predict earthquakes, and for some other purposes, gravity measurements of high precision are necessary. To collect experiences and initiate new methods, a new special study group was established at the General Assembly of the IAG in Lucerne.

The SSG prepared a preliminary report on special gravity measurements based on 13 answers received to a circular sent out by the president of the SSG. This preliminary report was supplemented by 9 new pieces of information on a new circular sent out in February 1971.

.....

3 - THE WORKING OF THE SSG

The special Study groups of the IAG usually do nothing else but collect experiences and ideas from the specific field appointed for the SSG, and circulate its report. Other functions would be possible only if sponsored by the IAG or some other institutions. So far this SSG has only reported on the studies made by various institutes. Most of the members have considered that "there is no basis for continuing unless it undertakes to co-ordinate a specific project".

The following reports on special accurate gravity measurements in various countries and some new instrumental or measuring technique ideas.

4 - SPECIAL ACCURATE GRAVITY MEASUREMENTS

.....

General remarks

The following sources of disturbances in gravity measurements have been commonly given.

...

1. Vibration during transport or ground vibration can cause serious inaccuracies in the precise measurements (A.C. HAMILTON, B.C BARLOW, W. BULLERWELL, M. BONATZ, T. MURPHY, E. GROten, etc.).
2. Rapid or slow temperature changes have an effect on reading and scale (A. KIVINIEMI, Cl. ELSTNER, R. SIGL, E. GROten).
3. Mechanical shocks, also clamping the instrument, cause changes in readings (J.G. TANNER, R. BREIN).
4. In addition to air pressure influence on the gravimeter, the mass of air above the station and the deformation of the earth's crust caused by pressure changes must be taken into consideration in high precision accuracy measurements (A. KIVINIEMI, R. BREIN, R. LECOLAZET).
5. The variation of ground water level can cause larger changes than is generally appreciated (G. BARTA, W. BULLERWELL).
6. The selection of sites for high precision measurements is very important. The gravity differences to be measured should be as small as possible.

5 - NEW CONSTRUCTIONS

In the Institut für Angewandte Geodäsie in Frankfurt a new gravimeter for accurate measurement of small gravity differences has been constructed and a testing instrument built (4). This consists of a vertical spring balance. The reading device is a capacitance bridge, the reading accuracy of which corresponds to $\pm 1 \mu\text{Gal}$. An electromagnetic spring is situated below the capacitor. The current in the electric coil is changed until the indication of the capacitor bridge is zero. The electric current in the coil is the measure of gravity difference. The elimination of seismic perturbations is carried out by air damping of the capacitor device and by electronic integration.

A continuously recording, drift-free cryogenic gravimeter with very high sensitivity is under development at Stanislaus State College, in California, for AFCRL (12). A niobium hollow sphere floats freely in a properly adjusted magnetic field and field gradient produced with niobium zirconium coils in super-conduction state at $4^{\circ}2 \text{ K}$. The motion of the ball induces magnetic flux changes, which are recorded. The theoretical accuracy of this instrument is 10^{-8} Gal . This instrument is suitable for studying :

1. The secular variation of gravity
2. The variation of station positions with respect to the centre of mass or the variation of the Newtonian constant "G".
3. The correlation between short and long period variations and the angular velocity of the Earth's rotation.
4. The dynamic processes occurring inside the Earth.

A dynamic gravity gradient gradiometer has been developed by Dr. R. FORWARD of the Hughes Research Laboratories (U.S.A.) (13).

The basic principle of the rotating gradient sensor is as follows: If a system of proof masses is rotated in the static gravitational field of an object, the gravitational force gradient of this field will induce dynamic forces on the proof masses with a frequency which is twice the rotation frequency of the system; at the same time, inertial effects, caused by accelerations of the proof mass mounting, will induce forces with a frequency at the rotational frequency. The gravitational gradient can therefore be determined independently of inertial forces, also if the proof mass system is in free fall.

The Laboratory prototype has reached an accuracy of ± 1 E.U. at an integration time of 10 sec.

The instrument will have applications in airborne, marine and satellite gravimetry and navigation.

Dr. M. BONATZ, at the Institut für Theoretische Geodesie der Universität Bonn, has constructed an astatized lever arm balance to measure small vertical differences in gravity (1). The instrument has been tested on the vertical calibration base of the University of Bonn and an accuracy of ± 0.01 mGal with one measurement has been achieved.

Prof. GROTHEN of the Technische Hochschule, Darmstadt, has used a LaCoste-Romberg gravimeter with a capacitance-read-out-system with a strip chart recorder to study vertical gradient. μ Gal accuracy has been obtained. Prof. GROTHEN also proposes to remeasure the Gravitational Constant using gravimeters and heavy masses which he has used for calibration of gravimeter (6).

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TABLES FOR CONDENSATION REDUCTION

J.C. BHATTACHARJI

Survey of India, Dehra Dun

Summary

By making use of the rigorous formula as already developed by the author (BHATTACHARJI 1969) for the determination of condensation reduction on the basis of spherical hypothesis for the mathematical form of the earth, suitable reduction tables are worked out in the form of fundamental and special tables easily adaptable for general use at condensation surfaces of any elevations including those of the Earth model and the geoid.

INTRODUCTION

One of the most important geodetic applications of gravity anomalies is the determination of the shape of the geoid. By making use of Green's theorem, one can ultimately obtain the form of the surface parallel to the geoid (Ellipsoid) on having known its gravity anomaly field provided all terrestrial masses external to the surface are imagined as being transferred either to infinity or to within the surface (Heiskanen and Vening Meinesz 1958). But from practical point of view it becomes essential that the intended mass-transfers should cause as little change in the shape of the surface as possible. Hence Stokes and later Helmert introduced the condensation reduction for the purpose. In this reduction all terrestrial masses external to the geoidal surface are condensed into a surface layer of infinitesimal thickness lying on the geoid itself. The ideal geoid after condensation is found to differ from the actual geoid by only 3 metres in extreme cases and perhaps 1 metre on the average (Lambert 1930). But in the case of the Earth model surface, the entire mass-transfers involved being very limited, the corresponding intermediate surface or the co-geop similar to the ideal geoid in the case of the geoid, does not appreciably differ from the Earth model surface we seek to study. Now as a result of condensation into a surface layer either on the geoid or the Earth model or any other condensation surface, the value of gravity observed at a station on the physical surface of the earth is bound to undergo change by an amount equal to the gravity effect E at a point on the condensation surface vertically below or above the station of observation, due to the condensed topographic mass minus the gravity effect T at the station of observation, due to the same amount of the actual topographic mass lying above the condensation surface considered. This is taken as approximately equal to the "Gelandereduktion" or "Topographische Korrektion" of Borrass's reports on gravity (Lambert 1930), but is significantly in error in the oceanic region as well as in regions of rugged topography,

especially when the station of observation is at a lower elevation, as will be quite clear from the examples cited at the end. The procedure followed by Cassinis in his Fundamental Tables, is also quite approximate, as he seems to have avoided the rigorous treatment of the problem thinking it to be too much complicated for a practical solution (Cassinis and Dore 1937).

An attempt has therefore been made to make use of the rigorous formula as already developed by the author (Bhattacharji 1969) and reproduced here for reference, for a precise solution on the basis of spherical hypothesis for the mathematical form of the earth and then design suitable reduction tables easily adaptable for general use. The gravity effect T of the actual topographic mass requires to be computed by the usual methods (Bullard 1936, Cassinis and Dore 1937, and Bhattacharji 1963 & 1965); while the effect E of the condensed topography evaluated with the help of the fundamental or special tables for condensation reduction given in this publication.

DERIVATION OF CONDENSATION REDUCTION FORMULA

In Fig. 2, let be H_S the elevation of the station S , and H (positive upwards above the geoid) be the elevation of the zone on the Hayford system, having uniform density of unity throughout and bounded by the angular distances θ_1 and θ_2 so that

$$h = H - H_S$$

$$r = R + H_S$$

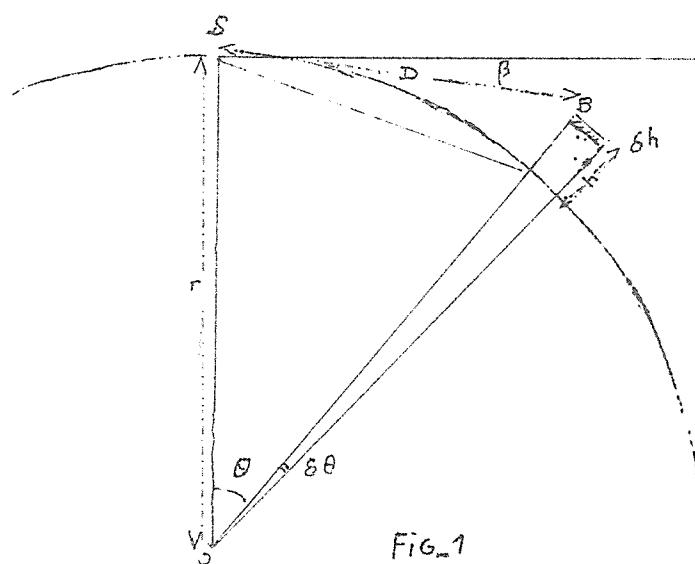


Fig. 1

where the mean radius R of the International Ellipsoid is equal to 6371.2 km.

Let an elementary mass dm at B (Fig. I) be at an angular distance θ and direct distance D from S.

Then the vertical component of the attractive effect dT at S of the elementary mass dm at B, acting upon a unit mass of one gramme at S, becomes :

$$dT = \frac{K dm}{D^2} \sin \beta$$

where $K = 6.67 \times 10^{-8}$ c.g.s., the constant of gravitation

$$dm = 2\pi(r + h)^2 \sin \theta \cdot d\theta \cdot dh$$

and β = the angle of depression / elevation of the straight distance from S to B, below / above the horizon at S. (Fig. I)

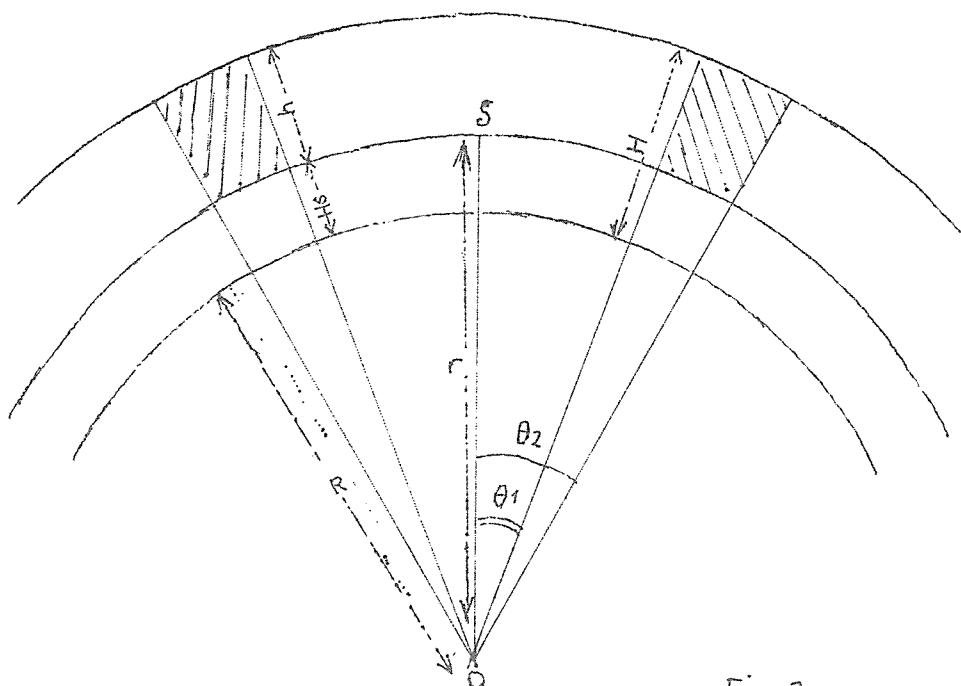


Fig-2

Now from Fig. I, we have,

$$\sin \beta = \frac{r - (r+h)\cos \theta}{D} = \frac{2(r+h)\sin^2 \theta/2 - h}{D}$$

$$\begin{aligned} dT &= \frac{2\pi K (r+h)^2 \sin \theta \cdot d\theta \cdot dh}{D^3} \left\{ 2(r+h) \sin^2 \frac{\theta}{2} - h \right\} \\ &= 2\pi K (1+y)^2 \left\{ 2(1+y)x^2 - y \right\} \frac{\sin \theta \cdot d\theta \cdot dh}{F^{3/2}} \end{aligned}$$

$$\text{where } x = \sin \frac{\theta}{2}, \quad y = \frac{h}{r}$$

$$F = \frac{D^2}{r^2} = \frac{1}{r^2} \left[r^2 + (r+h)^2 - 2r(r+h) \cos\theta \right] = 4x^2(1+y) + y^2$$

$$dh = rdy ; \quad \text{and } 1 - x^2 = \cos^2 \frac{\theta}{2}$$

$$\sin\theta d\theta = 2d(x^2)$$

thus $\frac{dT}{dT} = \frac{4\pi K(1+y)^2 [2x^2(1+y) - y]}{[4x^2(1+y) + y^2]^{3/2}} d(x^2) dh \quad (I)$

Now on integrating the above, and remembering that the gravity effect E of the actual topography condensed into a surface layer of thickness of 1 metre and of superficial density σ , on the condensation surface \bar{H} , is equivalent to $\sigma \left(\frac{dT}{dh} \right)$, where $h = \bar{H} - H_s$, \bar{H} is the elevation of the condensation surface, we have :

$$E = \sigma dT = \left[4\pi K \sigma (1+y)^2 \int \frac{[2x^2(1+y) - y]}{[4x^2(1+y) + y^2]^{3/2}} d(x^2) \right] \times 100 \text{ gal} \quad h = \bar{H} - H_s$$

or

$$E = 200 \pi k \sigma \left[\frac{(1+y)^2 (y+2x^2)}{[4x^2(1+y) + y^2]^{1/2}} \right]_{x_1}^{x_2}, \quad h = \bar{H} - H_s \quad (2)$$

where x_1 and x_2 denote the sines of half the angular distances of the inner and outer radii of the considered zone and superficial density σ of the condensed layer as given by Cassini's is

$$- \int_1 (H_1 - \bar{H}) \left(1 + \frac{H_1 - \bar{H}}{R} \right) - \frac{20}{3} \frac{\bar{H}^2}{R^2} + \dots$$

in the case of the continental zone of density σ_1 and elevation H_1 (in metres)

$$- \text{ and } \int_1 (H_2 - \bar{H}) \left(1 + \frac{H_2 - \bar{H}}{R} \right) - \frac{20}{3} \frac{\bar{H}^2}{R^2} + \dots = \int_2 H_2 \left(1 + \frac{H_2 - 2\bar{H}}{R} \right) - 5 \frac{\bar{H}^2}{R^2} + \dots$$

in the case of the oceanic zone of density σ_2 and depth H_2 .

$$\text{or } \int_1 (H_1 - \bar{H}) \left(1 + \frac{H_1 - \bar{H}}{R} \right)$$

in the case of the condensation surface at the Earth model surface

$$- \text{ and simply } \int_1 H_1 \left(1 + \frac{H_1}{R} \right)$$

in the case of the condensation surface at the geoid, the depth of the oceanic zone being completely ignored, in the first case due to all depths of the oceanic zone being converted into negative heights by the relation : $H_1 = -0.615 H_2$ while smoothening the Earth model surface, and in the second case due to the complete absence of any terrestrial masses lying above the geoid.

FUNDAMENTAL TABLES FOR REDUCING CONDENSATION EFFECTS AT CONDENSATION SURFACE OF ANY ELEVATION

The expression for condensation effect E as given in relation (2) has been evaluated zone by zone, first by assuming y to be equal to $\frac{h}{R}$ instead of $\frac{h}{R+H}$, for different values of h or $\bar{H} - H_s$ from -9000 m to $+9000$ m as given in the fundamental tables, and then by providing appropriate corrections known as corrections for station elevation due to y being assumed to be $\frac{h}{R}$ instead of $\frac{h}{R+H}$, for different values of H_s , and h from -9000 m to $+9000$ m, in the auxiliary tables given at the end of those fundamental tables where the corrections are considered appreciable. The tabular values have been given correct to 0.5×10^{-8} gal (m) in the case of the fundamental tables to $0.5 H_s \times 10^{-12}$ gal (m) in the case of the auxiliary tables, so that for all normal values of $H_1 - \bar{H}$ or $H_2 - \bar{H}$ and $\bar{H} - H_s$, and of densities δ_1 and δ_2 of the order of 3 and 1 respectively, the final results of the computed effect E of the condensed masses in any one zone may be of the precision of 0.1 mgal.

Example 1.

To find the condensation effect E of the continental zone F of elevation $H_1 = 9500$ and density $\delta_1 = 2.67$, at the station H_s , the elevation H_s and that of the condensation surface \bar{H} being 700 m and 500 m respectively.

We have

$$H_1 - \bar{H} = +9000 \text{ m}$$

$$T = 2.67 \times 9000 \left(1 + \frac{9000}{6371000} + \dots \right) = + 24064$$

$$h = \bar{H} - H_s = -200 \text{ m}$$

and from the fundamental tables,

(*) Tabular values 457 stands for 456.5, 395 for 394.5,
0.0001 H_s for 0.00005 H_s etc....

$$dT = +283 \times 10^{-8} \text{ gal} \quad (\#)$$

Hence the required condensation effect E

$$E = +283 \times 24064 \times 10^{-8} \text{ gal} = +68.0 \text{ mgal}$$

N.B. The same according to Cassinisi's method = +68.8 mgal

Example 2.

To find the condensation effect E of the continental zone G of elevation $H_1 = 8500 \text{ m}$ and density $\delta_1 = 2.67$, at the station H_s the elevation H_s and that of the condensation surface \bar{H} being 0 m and -500 m respectively.

We have,

$$H_1 - \bar{H} = +9000 \text{ m}$$

$$\bar{\sigma} = 2.67 \times 9000 \left(1 + \frac{9000}{6371000} + \dots\right) = +24064$$

$$h = \bar{H} - H_s = -500 \text{ m}$$

and from the fundamental tables,

$$dT = +305 \times 10^{-8} \text{ gal} \quad (\#)$$

Hence the required condensation effect E

$$= +305 \times 24064 \times 10^{-8} \text{ gal} = +73.4 \text{ mgal}$$

N.B. The same according to Cassinisi's method = +72.2 mgal

Example 3.

To find the condensation effect E of the oceanic zone N of depth $H_2 = -2000 \text{ m}$ and density $\delta_2 = 1.027$, the elevation of the station I_s and that of the condensation surface \bar{H} being 500 m and -7000 m respectively.

We have,

$$H_2 - \bar{H} = +5000 \text{ m}$$

$$\bar{\sigma} = 2.67 \times 5000 \left(1 + \frac{5000}{6371000} + \dots\right) + 1.027 \times 2000 \left(1 + \frac{12000}{6371000} + \dots\right)$$

$$= +15418$$

$$h = \bar{H} - H_s = -7500 \text{ m}$$

and from the fundamental tables,

$$dT = +227 \times 10^{-8} \text{ gal} \quad (\#)$$

Hence the required condensation effect E

$$= +227 \times 15418 \times 10^{-8} \text{ gal} = +34.9 \text{ mgal}$$

N.B. The same according to Cassinisi's method = +35.5 mgal

Correction to $dT = 0$

SPECIAL TABLES FOR REDUCING CONDENSATION EFFECTS AT EARTH MODEL AND GEOIDAL SURFACES

Use has been made of the fundamental tables to design 3 special tables for the computation of condensation effects of terrestrial masses at the level of the Earth model surface as well as of the geoid.

Table 1 has been designed for the purpose of obtaining the condensation effect E for the station elevation differences $h_s - H_s$ and the earth's topography $h - H > 0$ (topographic masses of heights for negative values of $h - H$ lying below the condensation surface H , negative values of $h - H$ to be taken as zero) condensed at the elevation H of the Earth model surface for zones H to 12, that for zones A to G and 11 to 1 being negligible.

Table 2 has been designed for the purpose of obtaining the condensation effect E for the terrain of the earth's topography $h - h_s$ (topographic masses of heights for negative values of h lying below the condensation surface at the geoid, negative values of h to be taken as zero *) condensed at the condensation surface at the geoid upto zone O_2 .

Table 3 has been designed for the purpose of reducing the topographic condensation effects $E-T$ for the terrain of the earth's topography $h - h_s$ (negative values of h to be taken as zero as before *) condensed at the condensation surface on the geoid for zones 18 to 12, that for zones 11 to 1 being negligible as in the case of Table 1. As far as the topographic condensation effects $E-T$ for the earth's plateaux of thickness h_s and density 2.67, considered from zone A to O_2 and also between zones 18 and 12 in the case of the condensation surface at the geoid are concerned, these are comparatively small in magnitude and therefore their combined effects have been given at the end of Tables 2 and 3, as corrections for station elevation for zones A to O_2 and 18 to 12 respectively, the corresponding corrections for zones 11 to 1 being negligible.

Apart from that, the gravity effect T due to the earth's entire topography lying above the Earth model surface in the case of the condensation surface considered at the Earth model, works out as $T_1 = \underline{O} - (\bar{O} - \bar{T})$ (Bhattacharji 1965),

where

$$T_1 = (A + 5.6 \times 10^{-3}) (h_s - H_s) \text{ or } 0.1174 (h_s - H_s),$$

which is easily computable by simple multiplications only.

\underline{O} = Terrain effect of the earth's topography $h - h_s$ for zones A to O_2 , which can be easily read off from the table given by the author in his previous paper (Bhattacharji, 1963, pp. 26-27), and

....

* In this case, the terrain effect \underline{O} or \bar{O} for elevations h is also to be subtracted from the computed results.

$(\bar{O} - \bar{\Omega})$ = Terrain effect \bar{O} of the earth's topography $h-h_s$ for zones 18 to 12 minus terrain effect of the topography of the Earth model $H-H_s$ for zones 18 to 12, which can be easily read off from the table given by the author in his previous paper (Bhattacharji 1965, pp. 24-25).

Thus the topographic effects as obtained above, when subtracted from the corresponding condensation effects E as read off from Table 1 and Tables 2 & 3 respectively, will furnish the topographic condensation effects $E-T$ for the entire earth in the case of the Earth model surface and the geoidal surface separately.

The tabular values have been given correct to 0.5×10^{-5} gal in the case of the main tables and to 1.0×10^{-4} gal in the case of the corrections for station elevations, so that the final results of the computed effects may be of the precision of about 0.1 mgal.

Example 1.

To find the condensation effect E of the continental zone K of elevation $h = 6000$ m and density 2.67, the elevation of the station H_s and that of the condensation surface H at the Earth model being 4000 m and 3000 m respectively.

We have :

$$h - H = + 3000 \text{ m}$$

$$H - H_s = - 1000 \text{ m}$$

and from Table 1,

$$E = + 0.47 \times 20 = + 9.3 \text{ mgal}$$

N.B. The same according to Cassinini's method = + 9.6 mgal. (≡)

Example 2.

To find the condensation effect E of the oceanic zone K of depth $d = - 4000$ m and density 1.027, the elevation of the station H_s and that of the condensation surface H at the Earth model being - 4000 m and - 5000 m respectively.

We have :

$$h = -0.615 \times 4000 \text{ m} = -2460 \text{ m}$$

$$h - H = -2460 \text{ m} + 5000 \text{ m} = +2540$$

$$H - H_s = -1000 \text{ m}$$

....

(≡) See p.T - II the results obtained by Cassinini's method and Lambert's formula.

and from Table 1,

$$E = +0.40 \times 20 = +7.9 \text{ mgal} \quad (\text{**})$$

Example 3.

To find the topographic-condensation effect E-T of the continental zones A to O_2 of elevations $h = 1000 \text{ m}$ for zones A to K and 5000 m for zones L to O_2 , and density 2.67, at the geoidal surface, the elevation of the station h_s being 1000 m .

we have

$$h - h_s = 0 \text{ for zones A to K and } +4000 \text{ m for zones L to } O_2.$$

From Table 2,

$$\begin{aligned} E &= +0.015 \times 24 + 0.075 \times 14 + 0.09 \times 16 + 0.04 \times 28 + 0.04 \times 28 \\ &= +5.1 \text{ mgal.} \end{aligned}$$

Correction for station elevation = $+0.4 \text{ mgal}$ and from the table of terrain effect \underline{O} (Bhattacharji 1963, pp. 26-27),

$$\underline{O} = -36.5 \text{ mgal}$$

Hence the required topographic-condensation effect E-T

$$= +5.1 + 0.4 + 36.5 = +42.0 \text{ mgal.} \quad (\text{**})$$

Example 4.

Same as Example 3 except that the elevations of the station h_s and of the zones A to K = 6000 m instead of 1000 m .

We have :

$$h - h_s = 0 \text{ for zones A to K, and } -1000 \text{ m for zones L to } O_2.$$

From Table 2 :

$$\begin{aligned} E &= -0.005 \times 24 - 0.02 \times 14 - 0.02 \times 16 - 0.01 \times 28 - 0.01 \times 28 \\ &= -1.3 \text{ mgal.} \end{aligned}$$

Correction for station elevation = $+12.7 \text{ mgal}$ and from the table of terrain effect \underline{O} (Bhattacharji 1963, pp. 26-27).

$$\underline{O} = -3.9 \text{ mgal}$$

Hence the required topographic-condensation effect E-T

$$= -1.3 + 12.7 + 3.9 = +15.3 \text{ mgal.} \quad (\text{**})$$

(**) See remark p. T - 8

Example 5.

To find the topographic-condensation effect E-T of zones A to O_2 of elevations $h = 1500$ m for zones A to M, and $d = 2000$ m for zones N to O_2 , and densities $\delta_1 = 2.67$ and $\delta_2 = 1.027$, at the geoidal surface, the elevation of the station h_s being 1500 m.

We have :

$$h - h_s = 0 \text{ for zones A to M, and } -1500 \text{ m for zones N to } O_2, \\ h \text{ being negative}$$

From Table 2,

$$\begin{aligned} E &= -0.035 \times 16 - 0.015 \times 28 - 0.015 \times 28 \\ &= -1.4 \text{ mgal} \end{aligned}$$

Correction for station elevation = +0.8 mgal and from the table of terrain effect Ω (Bhattacharji 1963, pp. 26-27)

$$\Omega = -5.0 \text{ mgal}$$

Hence the required topographic-condensation effect E-T

$$= -1.4 + 0.8 + 5.0 = +4.4 \text{ mgal} \quad (\text{xx})$$

Example 6.

To find the topographic-condensation effect E-T of the continental zones 18 to 12 of elevations $h = 4000$ m and density 2.67 at the geoidal surface, the elevation of the station h_s being 1000 m.

We have, :

$$h - h_s = +3000 \text{ m}$$

From Table 3,

$$\begin{aligned} E-T &= +0.035 \times 10 + 0.03 \times 10 + 0.03 \times 10 + 0.03 \times 10 + 0.03 \times 10 \\ &\quad + 0.025 \times 16 + 0.025 \times 10 \\ &= +2.2 \text{ mgal} \end{aligned}$$

Correction for station elevation = -0.3 mgal.

Hence the required topographic-condensation effect E-T

$$= +2.2 - 0.3 = +1.9 \text{ mgal.} \quad (\text{xx})$$

(xx) See remark p. T - 8

Example 7.

Same as Example 6 except that the elevation of the station
 $h_s = 7000$ m instead of 1000 m.

We have :

$$h - h_s = -3000 \text{ m}$$

From Table 3,

$$\begin{aligned} E-T &= +0.035 \times 10 + 0.03 \times 10 + 0.03 \times 10 + 0.03 \times 10 \\ &\quad + 0.03 \times 10 + 0.025 \times 16 + 0.025 \times 10 \\ &= +2.2 \text{ mgal.} \end{aligned}$$

Correction for station elevation = - 12.1 mgal.

Hence the required topographic-condensation effect E-T

$$= +2.2 - 12.1 = -9.9 \text{ mgal} \quad (\#)$$

Abstract of Results

Examples	Results obtained by			Discrepancy	
	Suggested method (1)	Cassini's method (2)	Lambert's method (3)	(1)-(2)	(1)-(3)
	mgal	mgal	mgal	mgal	mgal
1 p.T-6	+ 68.0	+ 69.8	-	- 1.8	-
2 p.T-6	+ 73.4	+ 72.2	-	+ 1.2	-
3 p.T-6	+ 34.9	+ 35.5	-	- 0.6	-
1 p.T-8	+ 9.3	+ 9.6	-	- 0.3	-
2 p.T-9	+ 7.9	+ 8.1	-	- 0.2	-
3 p.T-9	+ 42.0	+ 42.9	+ 36.5	- 0.9	+ 5.5
4 p.T-9	+ 15.3	+ 19.2	+ 3.9	- 3.9	+ 11.4
5 p.T-10	+ 4.4	+ 5.7	+ 7.1	- 1.3	- 2.7
6 p.T-10	+ 1.9	+ 2.0	0	- 0.1	+ 1.9
7 p.T-11	- 9.9	- 9.9	0	0.0	- 9.9

CONCLUSION

From the abstract of results obtained by various methods as illustrated above, it becomes immediately clear that the reduction tables incorporated in this publication are not only more easily adaptable for general use than any of the existing methods, but also most essential for the precise determination of the condensation reductions carried out on different hypotheses having condensation surfaces of any elevations including those of the geoid and the Earth model, and as such, will prove useful in solving various delicate problems relating to reductions of condensation correction involved in physical geodesy.

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F U N D A M E N T A L T A B L E S

in units of 10^{-8} Gal

Z O N E A
0 -- 2m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
0	± 4181	18	± 26	36	± 7	70	± 2
1	2317	19	23	37	6	75	2
2	1228	20	21	38	6	80	2
3	704	21	19	39	6	85	1
4	443	22	17	40	5	90	1
5	300	23	16	41	5	95	1
6	215	24	15	42	5	100	1
7	162	25	14	43	5	105	1
8	125	26	13	44	5	110	1
9	100	27	12	45	4	115	1
10	82	28	11	46	4	120	1
11	68	29	10	47	4	125	1
12	57	30	10	48	4	130	1
13	49	31	9	49	4	135	1
14	42	32	8	50	4	140	1
15	37	33	8	55	3	145	1
16	33	34	7	60	3	150	1
17	29	35	7	65	2	200	0
18	± 26	36	± 7	70	± 2	9000	± 0

Z O N E B
2m - 68m

h_m	$dT \times 10^8$						
0	± 0	120	± 544	240	± 159	600	± 27
5	3684	125	509	245	153	650	23
10	3500	130	477	250	147	700	20
15	3252	135	448	255	142	750	17
20	2988	140	421	260	137	800	15
25	2732	145	396	265	132	850	14
30	2490	150	374	270	127	900	12
35	2266	155	353	275	123	950	11
40	2061	160	334	280	119	1000	10
45	1874	165	316	285	115	1200	7
50	1705	170	300	290	111	1400	5
55	1553	175	285	295	107	1600	4
60	1416	180	270	300	104	1800	3
65	1293	185	257	320	92	2000	3
70	1183	190	245	340	82	2200	2
75	1085	195	234	360	73	2400	2
80	997	200	223	380	66	2600	2
85	917	205	213	400	59	2800	1
90	846	210	204	420	54	3000	1
95	782	215	195	440	49	3500	1
100	725	220	187	460	45	4000	1
105	673	225	179	480	42	5000	1
110	626	230	172	500	38	6000	1
115	583	235	165	550	32	7000	0
120	± 544	240	± 159	600	± 27	9000	± 0

Z O N E C
68m - 230m
 $(c_1 + c_2)$

h_m	$dTx10^8$	h_m	$dTx10^3$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 1	- 1100	+ 87	- 500	+ 345	- 280	+ 834
7000	2	1050	89	490	358	275	854
6000	3	1000	97	430	370	270	874
5000	4	950	107	470	384	265	895
4000	7	900	119	460	398	260	916
3800	7	880	124	450	412	255	938
3600	8	860	129	440	428	250	960
3400	9	840	135	430	444	245	983
3200	10	820	142	420	461	240	1007
3000	11	800	148	410	480	235	1031
2800	13	780	155	400	499	230	1056
2600	15	760	163	390	519	225	1081
2400	18	740	171	380	540	220	1107
2200	21	720	180	370	563	215	1134
2000	25	700	190	360	587	210	1161
1900	28	680	200	350	612	205	1189
1800	31	660	212	340	638	200	1218
1700	35	640	224	330	667	195	1247
1600	39	620	237	320	696	190	1277
1500	44	600	251	310	728	185	1307
1400	51	580	267	300	761	180	1338
1300	59	560	284	295	779	175	1369
1200	68	540	303	290	797	170	1400
1150	74	520	323	285	815	165	1432
- 1100	+ 87	- 500	+ 345	- 280	+ 834	- 160	+ 1464

ZONE C

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 160	+ 1464	- 40	+ 1407	+ 80	- 1817	+ 200	- 1218
155	1496	35	1288	85	1820	205	1190
150	1528	30	1150	90	1817	210	1162
145	1560	25	994	95	1808	215	1134
140	1591	20	820	100	1795	220	1107
135	1622	15	630	105	1777	225	1081
130	1653	10	428	110	1757	230	1056
125	1680	- 5	+ 217	115	1733	235	1037
120	1708	0	0	120	1708	240	1007
115	1733	+ 5	- 216	125	1680	245	983
110	1757	10	428	130	1652	250	960
105	1777	15	630	135	1622	255	938
100	1795	20	820	140	1591	260	916
95	1808	25	993	145	1560	265	895
90	1817	30	1150	150	1528	270	874
85	1820	35	1288	155	1496	275	854
80	1817	40	1407	160	1464	280	834
75	1806	45	1508	165	1432	285	815
70	1786	50	1593	170	1400	290	797
65	1756	55	1661	175	1369	295	779
60	1715	60	1715	180	1338	300	762
55	1661	65	1756	185	1307	310	728
50	1593	70	1786	190	1277	320	697
45	1508	75	1806	195	1247	330	667
- 40	+ 1407	+ 80	- 1817	+ 200	- 1218	+ 340	- 639

Z O N E C

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
+ 340	- 639	+ 500	- 346	+ 820	- 142	+ 1800	- 31
350	612	520	323	840	135	1900	28
360	587	540	303	860	130	2000	25
370	563	560	284	880	124	2200	21
380	540	580	267	900	119	2400	18
390	519	600	251	950	107	2600	15
400	499	620	237	1000	97	2800	13
410	480	640	224	1050	89	3000	11
420	462	660	212	1100	81	3200	10
430	444	680	200	1150	74	3400	9
440	428	700	190	1200	68	3600	8
450	412	720	180	1300	59	3800	7
460	398	740	172	1400	51	4000	7
470	384	760	163	1500	44	5000	4
480	370	780	156	1600	39	6000	3
490	358	800	148	1700	35	7000	2
+ 500	- 346	+ 820	- 142	+ 1800	- 31	+ 9000	- 2

h	± 1000
Corrections to $dTx10^8$	$\mp 0.0001 H_S$

Z O N E D $(D_1 + D_2)$
230m - 590m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 8	- 2000	+ 144	- 1040	+ 447	- 500	+ 1098
8000	10	1950	157	1020	467	480	1135
7500	11	1800	158	1000	475	460	1172
7000	13	1850	166	980	490	440	1209
6500	15	1800	175	960	505	420	1246
6000	17	1750	184	940	521	400	1282
5500	20	1700	194	920	538	390	1299
5000	25	1650	205	909	555	380	1316
4800	27	1600	216	880	574	370	1335
4600	29	1550	229	860	593	360	1349
4400	32	1500	243	840	613	350	1364
4200	35	1450	257	820	633	340	1379
4000	38	1400	274	800	655	330	1393
3800	42	1350	291	780	677	320	1405
3600	47	1300	317	760	701	310	1417
3400	52	1280	319	740	725	300	1427
3200	59	1260	327	720	751	290	1435
3000	67	1240	336	700	777	280	1442
2900	71	1220	346	680	805	270	1447
2800	76	1200	355	660	833	260	1449
2700	82	1180	365	640	863	250	1449
2600	88	1160	376	620	893	240	1447
2500	95	1140	386	600	925	235	1445
2400	+ 102	1120	397	580	958	230	1442
2300	117	1100	409	560	992	225	1438
2200	120	1080	421	540	1026	220	1433
2100	131	1060	434	520	1062	215	1427
- 2000	144	- 1040	+ 447	- 500	+ 1098	- 210	1427

ZONE D

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 210	+ 1421	- 75	+ 771	+ 60	- 634	+ 195	- 1395
205	1413	70	727	65	681	200	1405
200	1405	65	681	70	727	205	1413
195	1395	60	634	75	771	210	1421
190	1385	55	586	80	814	215	1427
185	1373	50	537	85	855	220	1433
180	1360	45	486	90	895	225	1438
175	1346	40	435	95	934	230	1441
170	1331	35	383	100	971	235	1445
165	1314	30	330	105	1006	240	1447
160	1297	25	276	110	1040	250	1449
155	1278	20	221	115	1073	260	1449
150	1257	15	167	120	1103	270	1447
145	1235	10	111	125	1133	280	1442
140	1212	5	56	130	1161	290	1435
135	1187	0	0	135	1187	300	1427
130	1161	+ 5	- 56	140	1212	310	1417
125	1133	10	111	145	1235	320	1405
120	1104	15	166	150	1257	330	1393
115	1073	20	221	155	1277	340	1379
110	1040	25	276	160	1297	350	1364
105	1006	30	329	165	1314	360	1349
100	971	35	382	170	1337	370	1333
95	934	40	435	175	1346	380	1316
90	896	45	486	180	1360	390	1299
85	856	50	537	185	1373	400	1282
80	814	55	586	190	1385	420	1246
- 75	+ 771	+ 60	- 634	+ 195	- 1395	+ 440	- 1209

Z O N E D

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
+ 440	- 1209	+ 880	- 574	+ 1350	- 292	+ 2900	- 71
460	1172	900	556	1400	274	3000	67
480	1135	920	538	1450	258	3200	59
500	1098	940	522	1500	243	3400	52
520	1062	960	505	1550	229	3600	47
540	1027	980	490	1600	217	3800	42
560	992	1000	475	1650	205	4000	38
580	958	1020	461	1700	194	4200	35
600	925	1040	447	1750	184	4400	32
620	894	1060	434	1800	175	4600	29
640	863	1080	421	1850	166	4800	27
660	833	1100	409	1900	158	5000	25
680	805	1120	398	1950	151	5500	27
700	777	1140	387	2000	144	6000	17
720	751	1160	376	2100	132	6500	15
740	726	1180	365	2200	127	7000	13
760	701	1200	355	2300	117	7500	11
780	678	1220	346	2400	102	8000	10
800	655	1240	337	2500	95	+ 9000	- 8
820	634	1260	328	2600	88		
840	613	1280	319	2700	82		
860	593	1300	311	2800	76		
+ 880	- 574	+ 1350	- 292	+ 2900	- 71		

h	± 1000	± 2000
Corrections to $dTx10^8$	$\mp 0.0001 H_s$	$\mp 0.0001 H_s$

Z O N E E $(E_1 + E_2)$
 590m - 1280m

h_m	$dTx \cdot 10^8$	h_m	$dTx \cdot 10^8$	h_m	$dTx \cdot 10^8$	h_m	$dTx \cdot 10^8$
- 9000	+ 33	- 3000	+ 257	- 1450	+ 740	- 540	+ 1201
8500	37	2900	273	1400	769	520	1194
8000	41	2800	289	1350	799	500	1185
7500	47	2700	307	1300	830	490	1180
7000	54	2600	327	1250	862	480	1174
6500	62	2500	348	1200	895	470	1167
6000	72	2450	360	1150	928	460	1160
5800	77	2400	372	1100	962	450	1152
5600	82	2350	384	1050	996	440	1143
5400	88	2300	397	1000	1029	430	1134
5200	95	2250	411	950	1063	420	1124
5000	102	2200	425	900	1095	410	1114
4800	110	2150	440	850	1125	400	1102
4600	119	2100	456	800	1152	390	1090
4400	130	2050	473	780	1162	380	1077
4200	141	2000	490	760	1171	370	1063
4000	155	1950	508	740	1179	360	1049
3900	162	1900	527	720	1187	350	1033
3800	170	1850	546	700	1194	340	1017
3700	178	1800	567	680	1200	330	1000
3600	187	1750	589	660	1204	320	982
3500	197	1700	611	640	1207	310	963
3400	207	1650	635	620	1209	300	944
3300	218	1600	660	600	1210	290	923
3200	230	1550	685	580	1209	280	902
3100	243	1500	712	560	1206	270	879
- 3000	+ 257	- 1450	+ 740	- 540	+ 1201	- 260	+ 856

Z O N E E

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 260	+ 856	- 10	+ 39	+ 240	- 807	+ 490	- 1179
250	832	0	0	250	832	500	1185
240	807	+ 10	- 38	260	856	520	1194
230	781	20	77	270	879	540	1201
220	755	30	115	280	901	560	1206
210	727	40	153	290	923	580	1208
200	699	50	190	300	943	600	1210
190	670	60	228	310	963	620	1209
180	640	70	265	320	982	640	1207
170	609	80	302	330	1000	660	1204
160	578	90	338	340	1017	680	1200
150	545	100	374	350	1033	700	1194
140	512	110	409	360	1048	720	1187
130	479	120	444	370	1063	740	1180
120	445	130	478	380	1077	760	1171
110	410	140	512	390	1090	780	1162
100	374	150	545	400	1102	800	1152
90	339	160	577	410	1113	850	1125
80	302	170	609	420	1124	900	1095
70	265	180	639	430	1134	950	1063
60	228	190	669	440	1143	1000	1030
50	191	200	699	450	1152	1050	996
40	153	210	727	460	1160	1100	962
30	115	220	754	470	1167	1150	928
20	77	230	781	480	1173	1200	895
- 10	+ 39	+ 240	- 807	+ 490	- 1179	+ 1250	- 862

Z O N E E

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
+ 1250	- 862	+ 1950	- 508	+ 2800	- 290	+ 4400	- 130
1300	830	2000	490	2900	273	4600	120
1350	799	2050	473	3000	258	4800	111
1400	769	2100	457	3100	244	5000	102
1450	740	2150	441	3200	231	5200	95
1500	713	2200	426	3300	219	5400	89
1550	686	2250	412	3400	207	5600	83
1600	660	2300	398	3500	197	5800	77
1650	635	2350	385	3600	187	6000	72
1700	612	2400	372	3700	178	6500	62
1750	589	2450	360	3800	170	7000	54
1800	568	2500	349	3900	162	7500	47
1850	547	2600	328	4000	155	8000	42
1900	527	2700	308	4200	142	8500	37
+ 1950	- 508	+ 2800	- 290	+ 4400	- 130	+ 9000	- 33

h	$\pm 1000 \dots \pm 2000$	$\pm 3000 \dots \pm 5000$
Corrections to $dTx10^8$	$\mp 0.0001 H_s$	$\mp 0.0001 H_s$

Z O N E F $(F_1 + F_2)$

1280m--2290m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 88	- 5100	+ 242	- 2700	+ 591	- 1060	+ 913
8800	91	5000	250	2600	615	1040	910
8600	95	4900	258	2500	640	1020	907
8400	100	4800	267	2400	666	1000	903
8200	104	4700	276	2300	692	980	899
8000	109	4600	286	2200	719	960	895
7800	114	4500	296	2100	746	940	890
7600	120	4400	306	2000	773	920	884
7400	126	4300	318	1900	800	900	878
7200	132	4200	329	1800	826	880	871
7000	139	4100	341	1700	850	860	864
6800	147	4000	354	1650	862	840	857
6600	155	3900	368	1600	872	820	848
6400	164	3800	382	1550	882	800	839
6200	173	3700	397	1500	892	780	830
6000	183	3600	413	1450	900	760	820
5900	189	3500	429	1400	907	740	809
5800	194	3400	446	1350	913	720	798
5700	200	3300	464	1300	918	700	786
5600	206	3200	483	1250	920	680	774
5500	213	3100	503	1200	921	660	760
5400	220	3000	523	1150	920	640	747
5300	227	2900	545	1100	917	620	732
5200	234	2800	567	1080	915	600	717
- 5100	+ 242	- 2700	+ 591	- 1060	+ 913	- 580	+ 701

Z O N E - F

h_m	$dTx \cdot 10^8$	h_m	$dTx \cdot 10^8$	h_m	$dTx \cdot 10^8$	h_m	$dTx \cdot 10^8$
- 580	+ 701	- 100	+ 144	+ 380	- 507	+ 860	- 864
560	685	80	116	400	529	880	871
540	668	60	87	420	551	900	878
520	650	40	58	440	572	920	884
500	631	- 20	29	460	592	940	889
480	612	0	+ 1	480	612	960	894
460	593	+ 20	- 29	500	631	980	899
440	572	40	58	520	649	1000	903
420	551	60	86	540	667	1020	907
400	529	80	115	560	684	1040	910
380	507	100	144	580	701	1060	913
360	484	120	172	600	717	1080	915
340	461	140	200	620	732	1100	917
320	437	160	228	640	746	1120	919
300	413	180	255	660	760	1140	920
280	388	200	282	680	773	1160	921
260	362	220	309	700	786	1180	921
240	336	240	335	720	798	1200	921
220	310	260	361	740	809	1250	920
200	283	280	387	760	820	1300	918
180	256	300	412	780	830	1350	913
160	228	320	436	800	839	1400	907
140	201	340	460	820	848	1450	900
120	172	360	484	840	856	1500	892
- 100	+ 144	+ 380	- 507	+ 860	- 864	+ 1550	- 883

Z O N E F

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
+ 1550	- 883	+ 2700	- 591	+ 4300	- 318	+ 5900	- 189
1600	873	2800	568	4400	307	6000	184
1650	862	2900	546	4500	297	6200	174
1700	850	3000	524	4600	286	6400	164
1750	838	3100	503	4700	277	6600	155
1800	826	3200	484	4800	268	6800	147
1850	813	3300	465	4900	259	7000	140
1900	800	3400	447	5000	250	7200	133
1950	787	3500	430	5100	242	7400	127
2000	774	3600	413	5200	235	7600	121
2100	747	3700	398	5300	227	7800	115
2200	719	3800	383	5400	220	8000	110
2300	693	3900	369	5500	213	8200	105
2400	666	4000	355	5600	207	8400	100
2500	641	4100	342	5700	201	8600	96
2600	616	4200	330	5800	195	8800	92
+ 2700	- 591	+ 4300	- 318	+ 5900	- 189	+ 9000	- 88

h	± 1000	$\pm 2000 \dots 4000$	$\pm 5000 \dots \pm 9000$
Corrections to $dTx10^8$	$\pm 0.0001 H_s$	$\mp 0.0001 H_s$	$\mp 0.0001 H_s$

Z O N E G

2290m→3520m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 158	- 5600	+ 337	- 3200	+ 589	- 1300	+ 618
8800	165	5500	339	3100	601	1250	606
8600	171	5400	347	3000	613	1200	593
8400	178	5300	356	2900	624	1150	580
8200	185	5200	365	2800	635	1100	565
8000	193	5100	374	2700	646	1050	549
7800	201	5000	383	2600	655	1000	532
7600	210	4900	393	2500	664	900	495
7400	219	4800	403	2400	671	800	454
7200	229	4700	413	2300	678	700	408
7000	239	4600	423	2200	682	600	359
6900	244	4500	434	2100	686	500	305
6800	250	4400	445	2000	687	400	249
6700	255	4300	456	1900	686	300	189
6600	261	4200	467	1800	682	200	128
6500	267	4100	479	1750	679	- 100	64
6400	274	4000	491	1700	676	0	+ 7
6300	280	3900	503	1650	672	+ 100	- 64
6200	287	3800	515	1600	666	200	127
6100	293	3700	527	1550	661	300	188
6000	300	3600	540	1500	654	400	248
5900	308	3500	552	1450	646	500	304
5800	315	3400	564	1400	638	600	358
5700	323	3300	577	1350	628	700	408
- 5600	+ 337	- 3200	+ 589	- 1300	+ 618	+ 800	- 453

Z O N E G

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
+ 800	- 453	+ 2600	- 655	+ 4400	- 446	+ 6200	- 288
900	495	2700	646	4500	435	6300	281
1000	532	2800	636	4600	424	6400	274
1100	565	2900	625	4700	414	6500	268
1200	593	3000	613	4800	404	6600	262
1300	617	3100	602	4900	394	6700	256
1400	637	3200	590	5000	384	6800	251
1500	654	3300	577	5100	375	6900	245
1600	666	3400	565	5200	366	7000	240
1700	676	3500	553	5300	357	7200	229
1800	682	3600	540	5400	348	7400	220
1900	686	3700	528	5500	340	7600	211
2000	687	3800	516	5600	332	7800	202
2100	686	3900	503	5700	324	8000	194
2200	683	4000	491	5800	316	8200	186
2300	678	4100	480	5900	309	8400	179
2400	672	4200	468	6000	301	8600	172
2500	664	4300	457	6100	294	8800	165
+ 2600	- 655	+ 4400	- 446	+ 6200	- 288	+ 9000	- 159

h	± 1000	± 2000	$\pm 3000 \dots \pm 9000$
Corrections to $dTx10^8$	$\pm 0.0001 H_s$	0	$\mp 0.0001 H_s$

ZONE H

3520m-5240m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 287	- 5600	+ 488	- 3200	+ 635	- 800	+ 297
8800	290	5500	495	3100	636	700	263
8600	299	5400	503	3000	636	600	228
8400	309	5300	511	2900	636	500	192
8200	319	5200	518	2800	634	400	155
8000	330	5100	526	2700	631	300	117
7800	341	5000	534	2600	628	200	79
7600	352	4900	541	2500	623	- 100	40
7400	364	4800	549	2400	616	0	+ 7
7200	376	4700	556	2300	608	+ 100	- 39
7000	389	4600	563	2200	599	200	78
6900	395	4500	570	2100	589	300	116
6800	402	4400	578	2000	576	400	154
6700	409	4300	584	1900	563	500	197
6600	415	4200	591	1800	547	600	227
6500	422	4100	597	1700	530	700	262
6400	429	4000	603	1600	517	800	296
6300	436	3900	609	1500	490	900	328
6200	443	3800	614	1400	468	1000	359
6100	451	3700	619	1300	443	1100	389
6000	458	3600	623	1200	417	1200	417
5900	465	3500	627	1100	390	1300	443
5800	473	3400	631	1000	360	1400	467
5700	480	3300	633	900	329	1500	489
- 5600	+ 488	- 3200	+ 635	- 800	+ 297	+ 1600	- 510

Z O N E H

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
+ 1600	- 510	+ 3200	- 635	+ 4800	- 549	+ 6400	- 430
1700	529	3300	633	4900	542	6500	423
1800	546	3400	631	5000	534	6600	416
1900	562	3500	628	5100	527	6700	410
2000	576	3600	624	5200	519	6800	403
2100	588	3700	619	5300	512	6900	396
2200	599	3800	615	5400	504	7000	390
2300	608	3900	609	5500	496	7200	377
2400	616	4000	604	5600	489	7400	365
2500	622	4100	598	5700	481	7600	353
2600	627	4200	591	5800	474	7800	342
2700	631	4300	585	5900	466	8000	331
2800	634	4400	578	6000	459	8200	321
2900	636	4500	571	6100	452	8400	310
3000	636	4600	564	6200	444	8600	301
3100	636	4700	557	6300	437	8800	291
+ 3200	- 635	+ 4800	- 549	+ 6400	- 430	+ 9000	- 282

h	± 1000	± 2000	± 3000	$\pm 4000 \dots \pm 9000$
Corrections to $dTx10^8$	$\pm 0.0001 H_s$	$\mp 0.0001 H_s$	0	$\mp 0.0001 H_s$

ZONE I

5240m-8440m

h_m	$dTx \cdot 10^8$						
- 9000	+ 564	- 6600	+ 700	- 3200	+ 699	- 800	+ 238
8900	570	6500	705	3100	690	700	210
8800	576	6400	710	3000	679	600	181
8700	581	6300	715	2900	668	500	152
8600	587	6200	720	2800	656	400	122
8500	593	6100	724	2700	644	300	92
8400	599	6000	728	2600	630	200	62
8300	605	5800	736	2500	615	- 100	32
8200	610	5600	743	2400	600	0	+ 1
8100	616	5400	749	2300	584	+ 100	- 30
8000	622	5200	754	2200	566	200	60
7900	628	5000	757	2100	548	300	90
7800	634	4800	759	2000	529	400	120
7700	640	4600	760	1900	509	500	149
7600	645	4400	758	1800	489	600	179
7500	651	4200	754	1700	467	700	208
7400	657	4000	749	1600	444	800	236
7300	663	3900	745	1500	421	900	264
7200	668	3800	740	1400	397	1000	292
7100	674	3700	735	1300	372	1100	319
7000	679	3600	730	1200	347	1200	345
6900	685	3500	723	1100	321	1300	370
6800	690	3400	716	1000	294	1400	395
6700	695	3300	708	900	266	1500	419
- 6600	+ 700	- 3200	+ 699	- 800	+ 238	+ 1600	- 443

ZONE I

h_m	$dTx \times 10^8$						
+ 1600	- 443	+ 3200	- 698	+ 5600	- 744	+ 7400	- 658
1700	465	3300	707	5800	737	7500	653
1800	487	3400	715	6000	729	7600	647
1900	508	3500	722	6100	725	7700	641
2000	528	3600	729	6200	720	7800	635
2100	547	3700	735	6300	716	7900	630
2200	565	3800	740	6400	711	8000	624
2300	582	3900	744	6500	706	8100	618
2400	598	4000	748	6600	701	8200	612
2500	614	4200	754	6700	696	8300	606
2600	629	4400	758	6800	691	8400	601
2700	642	4600	759	6900	686	8500	595
2800	655	4800	759	7000	680	8600	589
2900	667	5000	757	7100	675	8700	583
3000	678	5200	754	7200	669	8800	577
3100	689	5400	749	7300	664	8900	572
+ 3200	- 698	+ 5600	- 744	+ 7400	- 658	+ 9000	- 566

h	$\pm 1000 \dots \pm 3000$	$\pm 4000 \dots \pm 5000$	$\pm 6000 \dots \pm 9000$
Corrections to $dTx \times 10^8$	$\pm 0.0001 H_s$	0	$\pm 0.0001 H_s$

ZONE J

8440m-12440m

h_m	$dTx \times 10^8$						
- 9000	+ 595	- 4800	+ 560	+ 0	+ 2	+ 4800	- 559
8800	599	4600	549	400	62	5000	568
8600	603	4400	537	800	125	5200	577
8400	607	4200	524	1200	285	5400	585
8200	609	4000	509	1400	215	5600	592
8000	611	3800	494	1600	243	5800	598
7800	613	3600	477	1800	271	6000	603
7600	615	3400	459	2000	298	6200	607
7400	615	3200	440	2200	324	6400	610
7200	616	3000	419	2400	349	6600	613
7000	616	2800	398	2600	373	6800	614
6800	615	2600	375	2800	396	7000	615
6600	613	2400	351	3000	417	7200	616
6400	611	2200	326	3200	438	7400	616
6200	607	2000	300	3400	457	7600	615
6000	603	1800	274	3600	475	7800	613
5800	599	1600	246	3800	492	8000	611
5600	593	1400	217	4000	508	8200	609
5400	586	1200	188	4200	522	8400	606
5200	578	800	127	4400	535	8600	603
5000	570	400	65	4600	548	8800	600
- 4800	+ 560	- 0	+ 2	+ 4800	- 559	+ 9000	596

h	± 1000	± 2000	\dots	± 5000	± 6000	\dots	± 8000	± 9000
Corrections to $dTx \times 10^8$	0	$\pm 0.0001 H_s$			0		$\pm 0.0001 H_s$	

Z O N E K
12440m-12800m

h_m	$dTx \cdot 10^8$						
- 9000	+ 653	- 5000	+ 492	+ 0	+ 2	+ 5000	- 489
8800	650	4800	478	400	- 44	5200	502
8600	647	4600	463	800	90	5400	515
8400	642	4400	448	1200	135	5600	527
8200	637	4200	433	1600	179	5800	539
8000	632	4000	416	2000	222	6000	550
7800	627	3800	400	2200	243	6200	561
7600	620	3600	382	2400	264	6400	571
7400	614	3400	364	2600	284	6600	580
7200	607	3200	346	2800	304	6800	589
7000	599	3000	327	3000	323	7000	597
6800	591	2800	308	3200	342	7200	605
6600	582	2600	288	3400	361	7400	612
6400	573	2400	268	3600	379	7600	619
6200	563	2200	247	3800	396	7800	625
6000	553	2000	226	4000	413	8000	630
5800	542	1600	183	4200	429	8200	636
5600	530	1200	139	4400	445	8400	640
5400	518	800	94	4600	460	8600	645
5200	505	400	48	4800	475	8800	648
- 5000	+ 492	- 0	+ 2	+ 5000	- 489	+ 9000	- 652

h	± 1000	$\pm 2000 \dots \pm 8000$
Corrections to $dTx \cdot 10^8$	0	$\pm 0.0001 H_s$

Z O N E L
18800m-38800m

h_m	$dTx \cdot 10^8$						
- 9000	+ 562	- 4500	+ 332	+ 0	+ 4	+ 4500	- 326
8500	543	4000	299	500	- 36	5000	358
8000	522	3500	265	1000	74	5500	388
7500	499	3000	230	1500	112	6000	417
7000	475	2500	193	2000	150	6500	444
6500	450	2000	156	2500	187	7000	470
6000	422	1500	119	3000	223	7500	495
5500	394	1000	81	3500	259	8000	517
5000	363	500	42	4000	293	8500	538
- 4500	+ 332	- 0	+ 4	+ 4500	- 326	+ 9000	- 558

h	$\pm 1000 \dots \pm 2000$	$\pm 3000 \dots \pm 9000$
Corrections to $dTx \cdot 10^8$	0	$\pm 0.0001 H_s$

Z O N E M

28800m-58800m

h_m	$dTx \cdot 10^8$						
- 9000	+ 625	- 4500	+ 337	+ 0	+ 10	+ 4500	- 318
8500	596	4000	302	500	- 27	5000	353
8000	566	3500	267	1000	65	5500	387
7500	535	3000	231	1500	102	6000	420
7000	504	2500	194	2000	138	6500	453
6500	472	2000	158	2500	175	7000	485
6000	439	1500	121	3000	211	7500	517
5500	405	1000	84	3500	247	8000	548
5000	372	500	47	4000	283	8500	578
- 4500	+ 337	- 0	+ 10	+ 4500	- 318	+ 9000	- 607

Z O N E N

58800m-99000m

h_m	$dTx \cdot 10^8$						
- 9000	+ 268	- 4000	+ 129	+ 0	+ 13	+ 4000	- 102
8000	241	3000	100	500	- 1	5000	131
7000	213	2000	71	1000	16	6000	159
6000	185	1000	42	2000	45	7000	187
5000	157	500	28	3000	74	8000	215
- 4000	+ 129	- 0	+ 13	+ 4000	- 102	+ 9000	- 242

h	± 1000	\dots	± 2000	± 3000	\dots	± 6000	± 7000	\dots	± 9000
Corrections to $dTx \cdot 10^8$	Z.M		0		$\pm 0.0001 H_s$		$\pm 0.0001 H_s$		$\pm 0.0001 H_s$
	Z.N		0		0		0		$\pm 0.0001 H_s$

Z O N E 01
99000m-132850m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 107	- 4000	+ 54	+ 0	+ 11	+ 4000	- 32
8000	97	3000	44	500	6	5000	43
7000	86	2000	33	1000	1	6000	54
6000	76	1000	22	2000	11	7000	64
5000	65	500	17	3000	21	8000	73
- 4000	+ 54	- 0	+ 11	+ 4000	- 32	+ 9000	- 85

Z O N E 02
132850m-166735m

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
- 9000	+ 69	- 4000	+ 37	+ 0	+ 11	+ 4000	- 15
8000	62	3000	37	500	8	5000	21
7000	56	2000	24	1000	5	6000	28
6000	50	1000	18	2000	2	7000	34
5000	43	500	15	3000	8	8000	40
- 4000	+ 37	- 0	+ 11	+ 4000	- 15	+ 9000	- 47

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$sTx10^8$	h_m	$dTx10^8$
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Z O N E 18

- 9000	+ 32	- 4000	+ 18	+ 0	+ 7	+ 4000	- 5
7000	27	3000	15	1000	4	5000	7
6000	24	2000	13	2000	2	6000	10
5000	21	1000	10	3000	2	7000	13
- 4000	+ 18	- 0	+ 7	+ 4000	- 5	+ 9000	- 18

Z O N E 17

- 9000	+ 32	- 4000	+ 19	+ 0	+ 9	+ 4000	- 3
7000	27	3000	17	1000	6	5000	5
6000	24	2000	14	2000	3	6000	8
5000	22	1000	11	3000	1	7000	10
- 4000	+ 19	- 0	+ 9	+ 4000	- 3	+ 9000	- 16

Z O N E 16

- 9000	+ 33	- 4000	+ 21	+ 0	+ 11	+ 4000	- 0
7000	28	3000	18	1000	8	5000	3
6000	26	2000	16	2000	6	6000	5
5000	23	1000	13	3000	3	7000	8
- 4000	+ 21	- 0	+ 11	+ 4000	0	+ 9000	- 13

h_m	$dTx \times 10^8$						
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Z O N E 115

- 9000	+ 35	- 4000	+ 23	+ 0	+ 14	+ 4000	+ 4
7000	31	3000	21	1000	11	5000	1
6000	28	2000	18	2000	9	6000	2
5000	26	1000	16	3000	6	7000	4
- 4000	+ 23	- 0	+ 14	+ 4000	+ 4	+ 9000	- 9

Z O N E 114

- 9000	+ 39	- 4000	+ 28	+ 0	+ 18	+ 4000	- 3
7000	35	3000	25	1000	16	5000	6
6000	32	2000	23	2000	- 13	6000	4
5000	30	1000	20	3000	11	7000	2
- 4000	+ 28	- 0	+ 18	+ 4000	+ 9	+ 9000	- 4

Z O N E 113

- 9000	+ 79	- 4000	+ 61	+ 0	+ 47	+ 4000	+ 32
7000	72	3000	57	1000	43	5000	29
6000	68	2000	54	2000	39	6000	25
5000	65	1000	50	3000	36	7000	21
- 4000	+ 61	- 0	+ 47	+ 4000	+ 32	+ 9000	+ 14

h_m	$dTx \times 10^8$						
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Z O N E 12

- 9000	+ 73	- 4000	+ 62	+ 0	+ 53	+ 4000	+ 45
7000	69	3000	60	1000	51	5000	43
6000	67	2000	58	2000	49	6000	40
5000	64	1000	56	3000	47	7000	38
- 4000	+ 62	- 0	+ 53	+ 4000	+ 45	+ 9000	+ 34

Z O N E 11

- 9000	+ 92	- 4000	+ 83	+ 0	+ 76	+ 4000	+ 69
7000	88	3000	81	1000	75	5000	68
6000	87	2000	80	2000	73	6000	66
5000	85	1000	78	3000	71	7000	64
- 4000	+ 83	- 0	+ 76	+ 4000	+ 69	+ 9000	+ 61

Z O N E 10

- 9000	+ 116	- 4000	+ 110	+ 0	+ 105	+ 4000	+ 100
7000	114	3000	109	1000	104	5000	99
6000	113	2000	108	2000	103	6000	97
5000	111	1000	106	3000	101	7000	96
- 4000	+ 110	- 0	+ 105	+ 4000	+ 100	+ 9000	+ 94

h_m	$dTx \cdot 10^8$						
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Z O N E 9

- 9000	+ 132	- 4000	+ 128	+ 0	+ 124	+ 4000	+ 121
7000	130	3000	127	1000	124	5000	120
6000	129	2000	126	2000	123	6000	120
5000	129	1000	125	3000	122	7000	119
- 4000	+ 128	- 0	+ 124	+ 4000	+ 121	+ 9000	+ 117

Z O N E 8

- 9000	+ 243	- 4000	+ 240	+ 0	+ 236	+ 4000	+ 233
7000	242	3000	239	1000	236	5000	232
6000	241	2000	238	2000	235	6000	232
5000	240	1000	237	3000	234	7000	231
- 4000	+ 240	- 0	+ 236	+ 4000	+ 233	+ 9000	+ 229

Z O N E 7

- 9000	+ 218	- 4000	+ 216	+ 0	+ 215	+ 4000	+ 214
7000	217	3000	216	1000	215	5000	213
6000	217	2000	216	2000	214	6000	213
5000	217	1000	215	3000	214	7000	212
- 4000	+ 216	- 0	+ 215	+ 4000	+ 214	+ 9000	+ 212

h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$	h_m	$dTx10^8$
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Z O N E 6

- 9000	+ 330	- 4000	+ 328	+ 0	+ 327	+ 4000	+ 326
7000	329	3000	328	1000	327	5000	326
6000	329	2000	328	2000	327	6000	325
5000	328	1000	327	3000	326	7000	325
- 4000	+ 328	- 0	+ 327	+ 4000	+ 326	+ 9000	+ 325

Z O N E 5

- 9000	+ 514	- 4000	+ 514	+ 0	+ 513	+ 4000	+ 512
7000	514	3000	513	1000	513	5000	512
6000	514	2000	513	2000	512	6000	512
5000	514	1000	513	3000	512	7000	512
- 4000	+ 514	- 0	+ 513	+ 4000	+ 512	+ 9000	+ 511

Z O N E 4

- 9000	+ 664	- 4000	+ 664	+ 0	+ 664	+ 4000	+ 663
7000	664	3000	664	1000	664	5000	663
6000	664	2000	664	2000	663	6000	663
5000	664	1000	664	3000	663	7000	663
- 4000	+ 664	- 0	+ 664	+ 4000	+ 663	+ 9000	+ 663

h_m	$dTx \cdot 10^8$						
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Z O N E 3

- 9000	+ 873	- 4000	+ 873	+ 0	+ 873	+ 4000	+ 873
7000	873	3000	873	1000	873	5000	873
6000	873	2000	873	2000	873	6000	873
5000	873	1000	873	3000	873	7000	874
- 4000	+ 873	- 0	+ 873	+ 4000	+ 873	+ 9000	+ 874

Z O N E 2

- 9000	+ 714	- 4000	+ 714	+ 0	+ 714	+ 4000	+ 715
7000	714	3000	714	1000	715	5000	715
6000	714	2000	714	2000	715	6000	715
5000	714	1000	714	3000	715	7000	715
- 4000	+ 714	- 0	+ 714	+ 4000	+ 715	+ 9000	+ 715

Z O N E 1

- 9000	+ 134	- 4000	+ 134	+ 0	+ 134	+ 4000	+ 134
7000	134	3000	134	1000	134	5000	134
6000	134	2000	134	2000	134	6000	134
5000	134	1000	134	3000	134	7000	134
- 4000	+ 134	- 0	+ 134	+ 4000	+ 134	+ 9000	+ 135

S P E C I A L T A B L E S

1 - 2 - 3

TABLE 1

(See p. T - 7)

EFFECT E for EARTH'S TOPOGRAPHY $h-H$
condensed at EARTH MODEL
COMPARTMENT^(a) of ZONES A to ₁^(c)

Unit = 0.01 mgal ; density : 2,67^(b)

- (a) The number of compartments n is given for each zone.
- (b) In calculating average heights, depths D (or d) are to be converted into heights H (or h) by the relation $H = -0,615 D$.
- (c) Effect for zone A is A ($h_F - H_F$) in which $A = 11.18/m$
Effect for zones B to G and zones 11 to 1 is negligible.

Difference of elevation $h - H$ in metres	ZONE H (n = 16)			ZONE I (n = 20)		
	$H - H_S$ in metres			$H - H_S$ in metres		
	+ 100	0	- 100	+ 200	0	- 200
0	0	0	0	0	0	0
100	- 1	0	+ 1	- 1	0	+ 1
200	- 2	0	+ 2	- 2	0	+ 2
300	- 3	0	+ 3	- 3	0	+ 3
400	- 4	0	+ 4	- 4	0	+ 4
500	- 5	0	+ 5	- 5	0	+ 5
600	- 6	0	+ 6	- 6	0	+ 6
700	- 7	0	+ 7	- 7	0	+ 7
800	- 8	0	+ 8	- 8	0	+ 8
900	- 9	0	+ 9	- 9	0	+ 9
1000	- 10	0	+ 10	- 10	0	+ 10
1200	- 12	0	+ 12	- 12	0	+ 12
1400	- 13	0	+ 13	- 13	0	+ 13
1600	- 14	0	+ 14	- 14	0	+ 14
1800	- 15	0	+ 15	- 15	0	+ 15
2000	- 16	0	+ 16	- 16	0	+ 16

Difference of elevation $h - H$ in metres	ZONE J ($n = 16$)			ZONE K ($n = 20$)		
	$H - H_s$ in metres			$H - H_s$ in metres		
	+ 500	0	- 500	+ 1000	0	- 1000
0	0	0	0	0	0	0
100	- 2	0	+ 2	- 2	0	+ 2
200	- 3	0	+ 3	- 3	0	+ 3
300	- 4	0	+ 4	- 5	0	+ 5
400	- 5	0	+ 6	- 6	0	+ 6
500	- 7	0	+ 7	- 8	0	+ 8
600	- 8	0	+ 8	- 9	0	+ 10
700	- 9	0	+ 10	- 11	0	+ 11
800	- 11	0	+ 11	- 12	0	+ 13
900	- 12	0	+ 12	- 14	+ 1	+ 14
1000	- 13	0	+ 14	- 15	+ 1	+ 16
1200	- 16	+ 1	+ 16	- 18	+ 1	+ 19
1400	- 18	+ 1	+ 15	- 21	+ 1	+ 22
1600	- 21	+ 1	+ 22	- 24	+ 1	+ 25
1800	- 24	+ 1	+ 24	- 27	+ 1	+ 28
2000	- 26	+ 1	+ 27	- 30	+ 1	+ 31
2200	- 29	+ 1	+ 30	- 33	+ 1	+ 34
2400	- 31	+ 1	+ 32	- 36	+ 1	+ 38
2600	- 34	+ 1	+ 35	- 39	+ 1	+ 41
2800	- 37	+ 1	+ 38	- 42	+ 1	+ 44
3000	- 39	+ 1	+ 40	- 45	+ 1	+ 47

Difference of elevation $h - H$ in metres	ZONE L ($n = 24$)			ZONE M ($n = 14$)		
	$H - H_s$ in metres			$H - H_s$ in metres		
	+ 1000	0	- 1000	+ 1000	0	- 1000
0	0	0	0	0	0	0
100	- 1	0	+ 1	- 1	0	+ 2
200	- 2	0	+ 2	- 3	+ 1	+ 3
300	- 3	0	+ 3	- 4	+ 1	+ 5
400	- 4	0	+ 4	- 5	+ 1	+ 7
500	- 4	0	+ 5	- 6	+ 1	+ 8
600	- 5	0	+ 6	- 8	+ 1	+ 10
700	- 6	+ 1	+ 7	- 9	+ 2	+ 11
800	- 7	+ 1	+ 7	- 10	+ 2	+ 13
900	- 8	+ 1	+ 8	- 11	+ 2	+ 15
1000	- 8	+ 1	+ 9	- 13	+ 2	+ 16
1200	- 10	+ 1	+ 11	- 15	+ 3	+ 19
1400	- 12	+ 1	+ 13	- 17	+ 3	+ 23
1600	- 13	+ 1	+ 15	- 20	+ 3	+ 26
1800	- 15	+ 1	+ 16	- 22	+ 4	+ 29
2000	- 17	+ 1	+ 18	- 25	+ 4	+ 32
2200	- 18	+ 1	+ 20	- 27	+ 4	+ 36
2400	- 20	+ 1	+ 22	- 30	+ 5	+ 39
2600	- 22	+ 1	+ 24	- 32	+ 5	+ 42
2800	- 23	+ 1	+ 25	- 35	+ 6	+ 45
3000	- 25	+ 1	+ 27	- 37	+ 6	+ 48

Difference of elevation $h - H$ in metres	ZONE N ($n = 16$)			ZONE O ₁ ($n = 28$)		
	$H - H_s$ in metres			$H - H_s$ in metres		
	+ 1000	0	- 1000	+ 2000	0	- 2000
0	0	0	0	0	0	0
500	- 2	+ 1	+ 4	- 1	+ 1	+ 2
1000	- 2	2	+ 7	- 1	+ 1	+ 3
1500	- 4	+ 4	+ 11	- 2	+ 2	+ 5
2000	- 5	+ 5	+ 14	- 2	+ 2	+ 6
2500	- 7	+ 6	+ 18	- 3	+ 3	+ 8
3000	- 8	+ 7	+ 21	- 3	+ 3	+ 10

Difference of elevation $h - H$ in metres	ZONE O ₂ ($n = 28$)		
	$H - H_s$ in metres		
	+ 2000	0	- 2000
0	0	0	0
500	0	+ 1	+ 1
1000	0	+ 1	+ 3
1500	0	+ 2	+ 4
2000	- 1	+ 2	+ 5
2500	- 1	+ 3	+ 6
3000	- 1	+ 3	+ 7

Difference of elevation $H - H_s$ in metres	ZONE 18 ($n = 10$)			ZONE 17 ($n = 10$)		
	$H - H_s$ in metres			$H - H_s$ in metres		
	+ 5000	0	- 5000	+ 4000	0	- 4000
0	0	0	0	0	0	0
200	- 1	+ 1	+ 1	0	+ 1	+ 1
400	- 1	+ 1	+ 2	0	+ 1	+ 2
600	- 1	+ 2	+ 4	- 1	+ 2	+ 3
800	- 2	+ 2	+ 5	- 1	+ 2	+ 4
1000	- 2	+ 2	+ 6	- 1	+ 2	+ 5
1200	- 3	+ 3	+ 7	- 1	+ 3	+ 6
1400	- 3	+ 3	+ 8	- 1	+ 3	+ 7
1600	- 3	+ 3	+ 9	- 1	+ 4	+ 8
1800	- 4	+ 4	+ 10	- 1	+ 4	+ 9
2000	- 4	+ 4	+ 11	- 1	+ 5	+ 10

Difference of elevation $H - H_s$ in metres	ZONE 16 ($n = 10$)			ZONE 15 ($n = 10$)		
	$H - H_s$ in metres			$H - H_s$ in metres		
	+ 5000	0	- 5000	+ 5000	0	- 5000
0	0	0	0	0	0	0
200	0	+ 1	+ 1	0	+ 1	+ 2
400	0	+ 1	+ 3	0	+ 2	+ 3
600	- 1	+ 2	+ 4	0	+ 2	+ 4
800	- 1	+ 2	+ 5	0	+ 3	+ 6
1000	- 1	+ 3	+ 6	+ 1	+ 4	+ 7
1200	- 1	+ 4	+ 8	+ 1	+ 5	+ 8
1400	- 1	+ 4	+ 9	+ 1	+ 5	+ 10
1600	- 1	+ 5	+ 10	+ 1	+ 6	+ 11
1800	- 1	+ 5	+ 11	+ 1	+ 7	+ 13
2000	- 1	+ 6	+ 13	+ 1	+ 7	+ 14

Difference of elevation $h - H$ in metres	ZONE 14 ($n = 10$)			ZONE 13 ($n = 16$)		
	$H - H_s$ in metres			$H - H_s$ in metres		
	+ 5000	0	- 5000	+ 5000	0	- 5000
0	0	0	0	0	0	0
200	+ 1	+ 1	+ 2	+ 1	+ 2	+ 2
400	+ 1	+ 2	+ 3	+ 2	+ 3	+ 5
600	+ 1	+ 3	+ 5	+ 3	+ 5	+ 7
800	+ 2	+ 4	+ 7	+ 4	+ 6	+ 9
1000	+ 2	+ 5	+ 8	+ 5	+ 8	+11
1200	+ 2	+ 6	+ 10	+ 6	+ 10	+13
1400	+ 3	+ 7	+ 11	+ 7	+ 11	+15
1600	+ 3	+ 8	+ 13	+ 8	+ 13	+17
1800	+ 3	+ 9	+ 14	+ 9	+ 14	+20
2000	+ 4	+ 10	+ 16	+ 10	+ 16	+22

Difference of elevation $h - H$ in metres	ZONE 12 ($n = 10$)		
	$H - H_s$ in metres		
	+ 5000	0	- 5000
0	0	0	0
200	+ 3	+ 3	+ 4
400	+ 5	+ 6	+ 7
600	+ 7	+ 9	+11
800	+ 9	+ 12	+14
1000	+12	+14	+17
1200	+14	+17	+21
1400	+16	+20	+24
1600	+18	+23	+28
1800	+21	+26	+31
2000	+23	+29	+34

T A B L E 2

(See p. T - 7)

EFFECT E for EARTH'S TOPOGRAPHY $h-h_s$
condensed at GEOID
COMPARTMENT (a) of ZONES A to O₂

Sign same as for $h-h_s$ (b)

Unit : 0,01 mGal; density : 2,67

- (a) For numbers of compartments in each zone, refer to the table 1
(b) h when negative is to be taken as zero

Differences of elevation $h - h_s$ in metres	Compartment of zones								
	A-H	I	J	K	L	M	N	O ₁	O ₂
0	0	0	0	0	0	0	0	0	0
500	0	0	0	0	0	1	1	1	1
1000	0	0	0	1	1	2	2	1	1
1500	0	0	1	1	1	3	4	2	2
2000	0	1	1	1	1	4	5	2	2
2500		1	1	1	1	5	6	3	3
3000			1	1	1	6	7	3	3
3500				1	2	7	8	4	4
4000					2	8	9	4	4
4500							10	5	5
5000							11	6	6
5500								6	6
6000								7	7

A= Elevation of station h_s in kilometersB= Correction for station elevation - Zones A to O₂

Unit: 0.1 mGal

A	0,5	1	1,5	2	2,5	3	3,5	4	4,5	5	5,5	6	6,5	7	7,5	8	8,5	9km
B	2	4	8	15	22	32	44	57	72	89	107	127	149	173	199	225	255	286mGal

TABLE 3 (See p. T - 7)

EFFECT E-T for EARTH'S TOPOGRAPHY $h - h_s$
 with condensation at GEOID
 COMPARTMENT (a) OF ZONES 18 to 1 (b)

Unit : 0.01 mGal : density : 2,67

- (a) For number of compartments in each zone refer to the Table 1
 (b) Effect E-T for zones 11 to 1 is negligible
 (c) h when negative is to be taken as zero

P = Station above compartment Q = Station below compartment.

Differences of elevations $ h - h_s $ (c) in metres	COMPARTMENT OF ZONES												
	18		17		16		15		14		13		12
	P	Q	P	Q	P	Q	P	Q	P	Q	P	Q	
0	0	0	0	0	0	0	0	0	0	0	0	0	0
500	0	0	0	0	0	0	0	0	0	0	0	0	0
1000	1	1	1	1	1	1	1	1	1	0	1	1	1
1500	1	1	1	1	1	1	1	1	1	1	1	1	1
2000	2	2	2	2	2	2	2	1	1	2	1	1	1
2500	3	3	2	2	2	2	2	2	2	2	2	2	2
3000	4	4	3	3	3	3	3	3	3	3	3	3	3
3500	5	5	5	5	4	4	4	4	4	4	4	4	4
4000	6	6	6	6	6	6	6	5	5	5	5	5	5
4500	8	8	7	7	7	7	7	7	7	7	6	6	6
5000	9	10	9	9	9	9	8	8	8	8	8	7	7
5500	11	12	11	11	10	10	10	10	10	10	9	9	9
6000	13	14	13	13	12	12	12	12	11	12	11	11	11

A = Elevation of station h_s in kilometersB = Correction for station elevation - Zones 18 to 12
 Unit : 0.1 mGal

A	0,5	1	1,5	2	2,5	3	3,5	4	4,5	5	5,5	6	6,5	7	7,5	8	8,5	9km
B	-1	-2	-5	-10	-16	-22	-31	-39	-50	-62	-75	-89	-105	-121	-139	-158	-179	-200 mGal

T A B L E D
BULLARD'S B CORRECTIONS

Unit : 0.1 mGal

Station elevation in metres	Correction
0	0
500	7
1000	11
1500	14
2000	16
2500	16
3000	14
3500	10
4000	5
4500	- 2
5000	- 11
5500	- 21
6000	- 33
6500	- 47
7000	- 62
7500	- 79
8000	- 98
8500	-118
9000	-140