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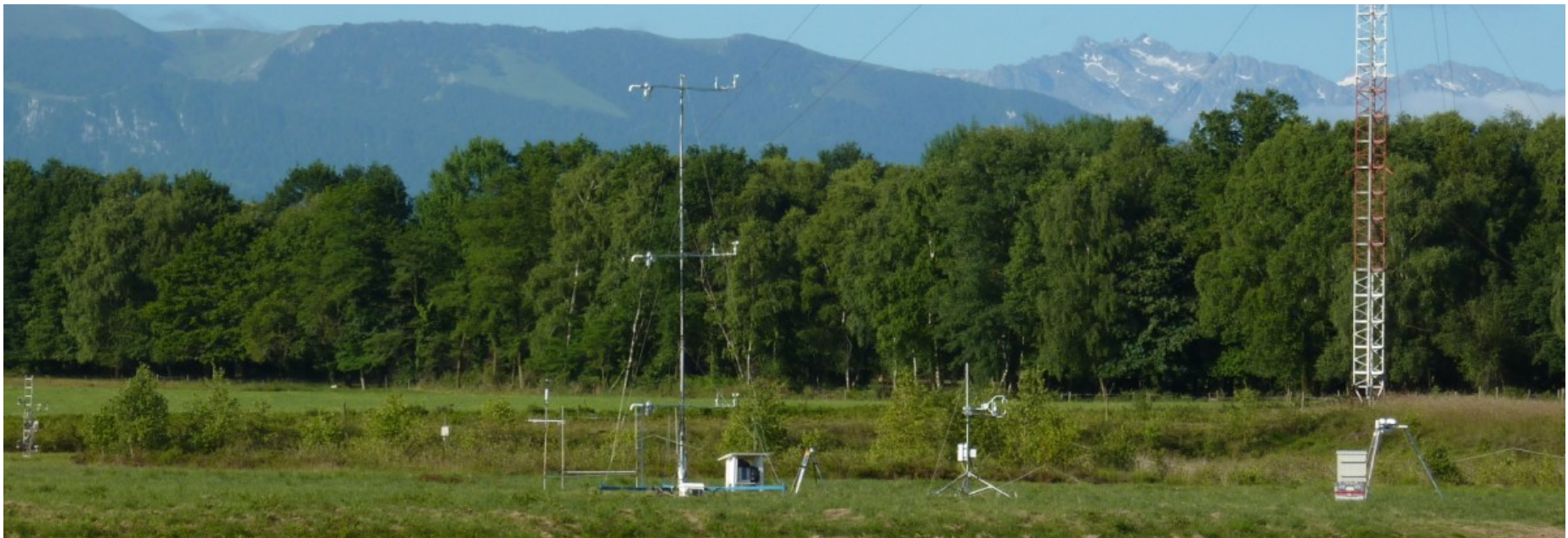
# Monin-Obukhov Similarity Theory in the 2-m column of thermocouples at the small-scale heterogeneity site

Maren Haid<sup>(1)</sup>, Daniel Martínez<sup>(1,2)</sup>, Jens Bange<sup>(1)</sup> and Joan Cuxart<sup>(2)</sup>

<sup>(1)</sup> *Center for Applied Geosciences, Eberhard Karls University Tübingen, Germany*

<sup>(2)</sup> *Group of Meteorology, University of the Balearic Islands, Mallorca, Spain*

[daniel.martinez@uni-tuebingen.de](mailto:daniel.martinez@uni-tuebingen.de)



# Introduction

**Objective:** Assessment of MOST in the lower surface layer at the SSH site of BLLAST campaign.

- Data:
  - Eight thermocouples within the first 2 m above ground level (AGL), Mallorca.
  - Eddy-covariance station at 1.95 m AGL, Bergen.
- Data requirements:
  - Steady state
    - \* Fair weather conditions
    - \* Excluding morning and evening transitions.
  - Horizontally homogeneous conditions: length scale required for an undisturbed fetch (Wyngaard, 2010):

$$L_x \gg \frac{\bar{u}}{u_*} z$$

## Introduction (II)



### Thermocouples:

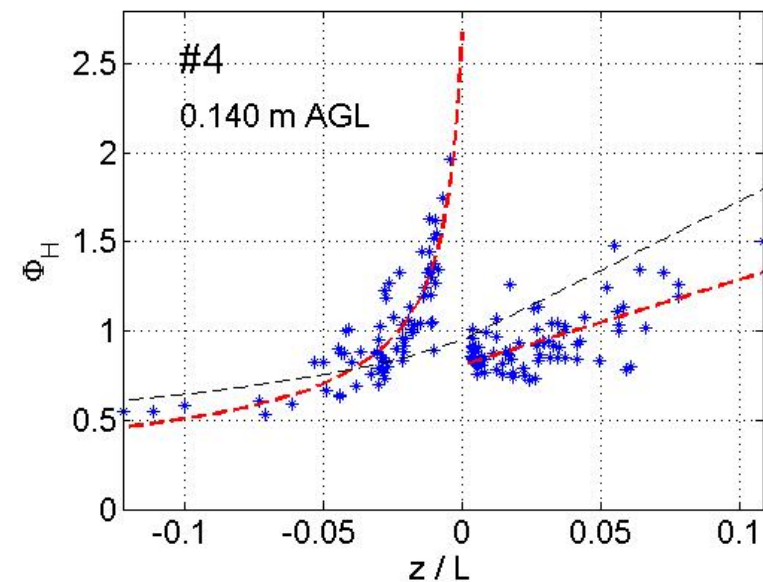
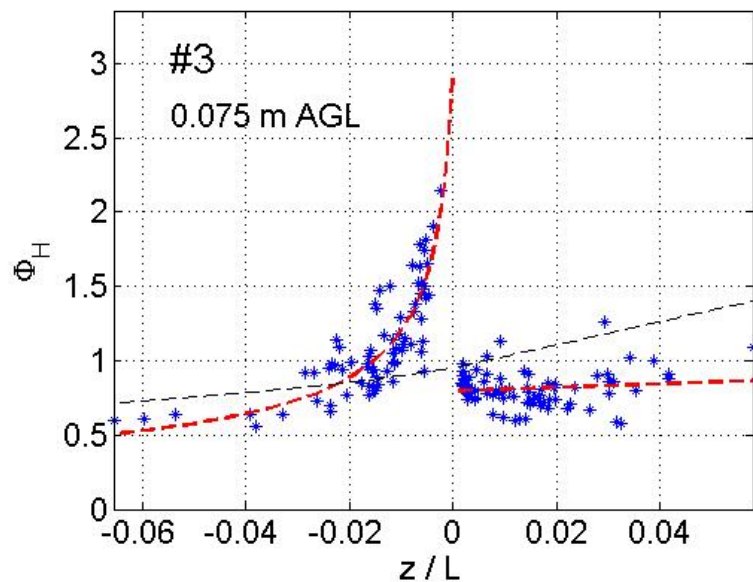
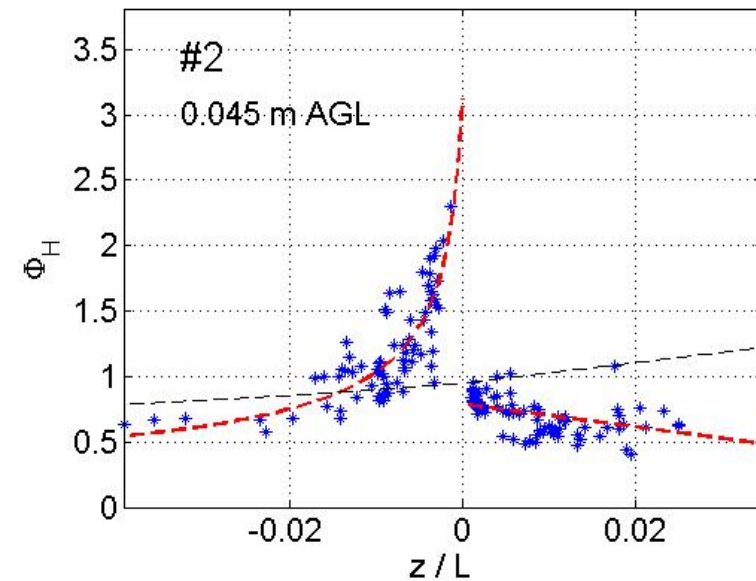
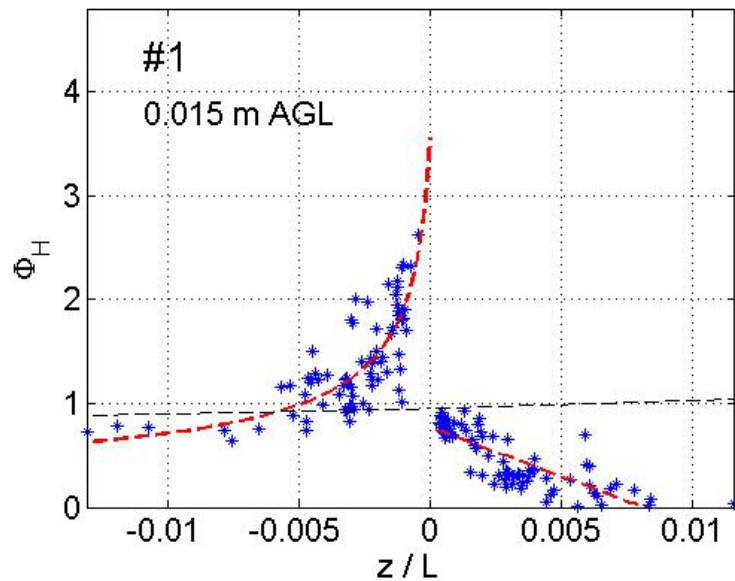
#	1	2	3	4	5	6	7	8
$z$ (m)	0.015	0.045	0.075	0.140	0.300	0.515	1.045	1.920
$z/z_0$	0.943	2.83	4.72	8.81	18.9	32.4	65.7	121

- For  $z_0 \simeq 1.59 \pm 0.89$  cm.
- MOST is valid for  $z/z_0 \gg 1 \Rightarrow$  We expect deviations for thermocouples #1–4.

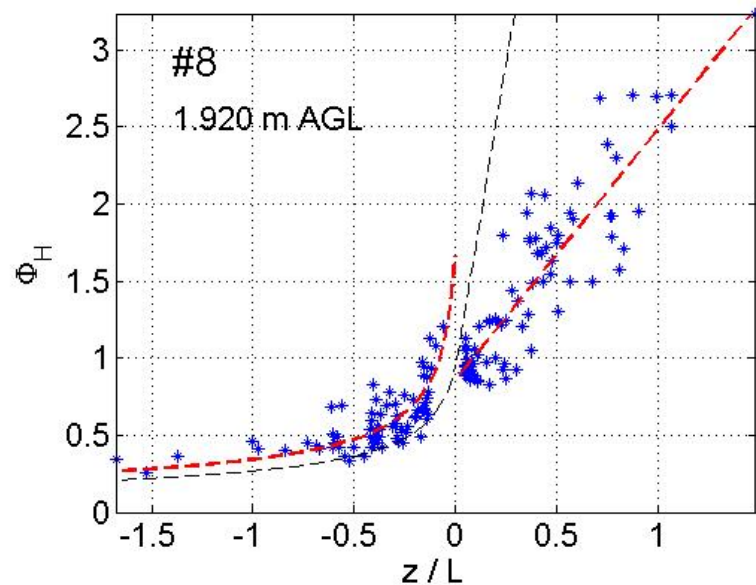
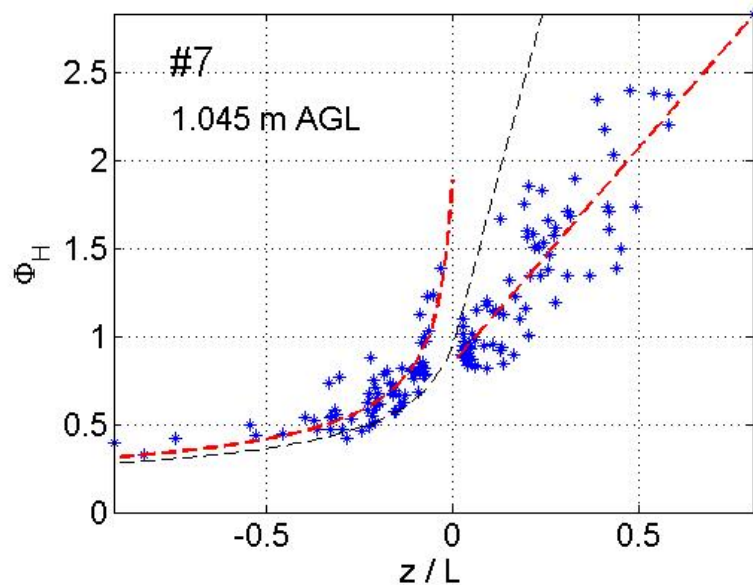
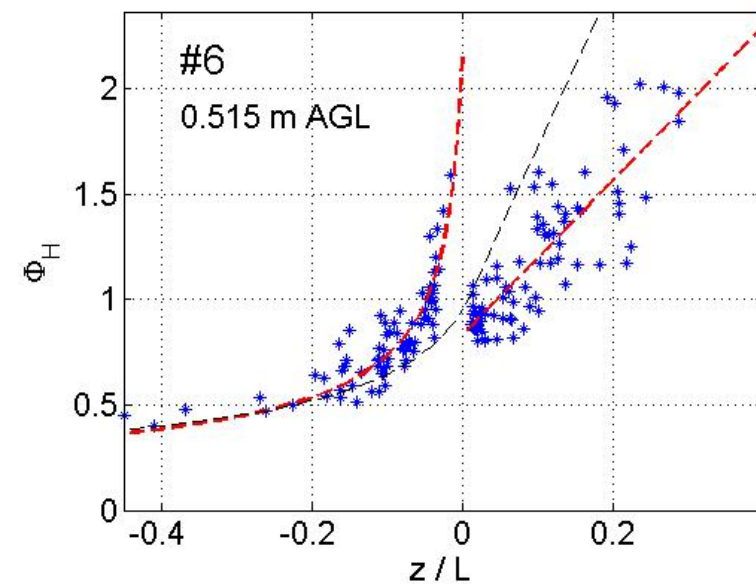
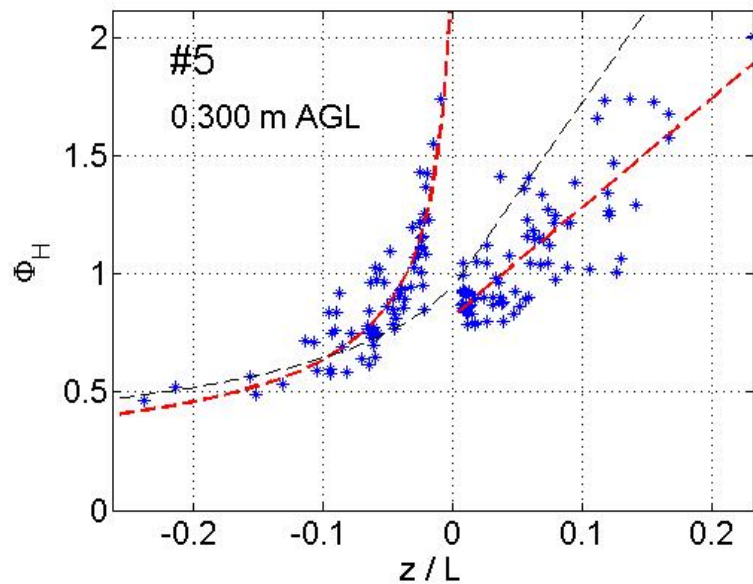
**Averaging time periods:** We select two different averaging time periods:

- 15 minutes
- 30 minutes

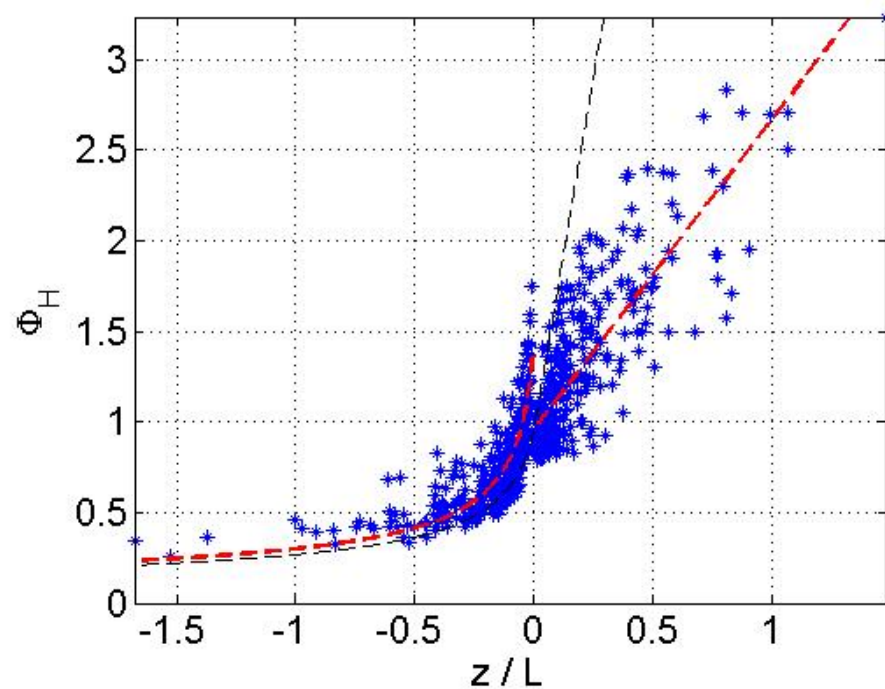
# Results (1a): MOST for each height



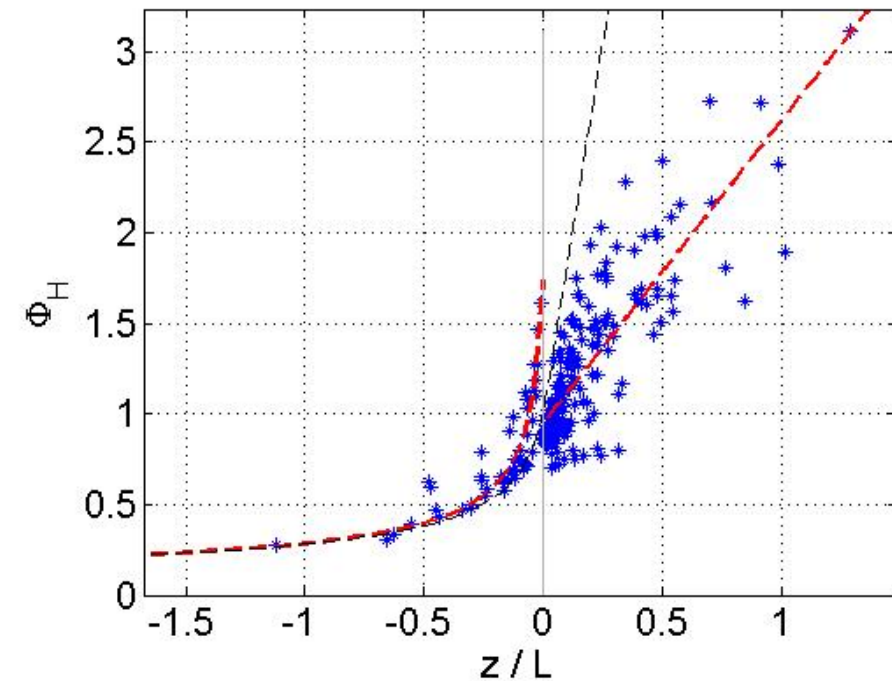
## Results (Ib): MOST for each height



## Results (II): MOST for the 4 uppermost heights



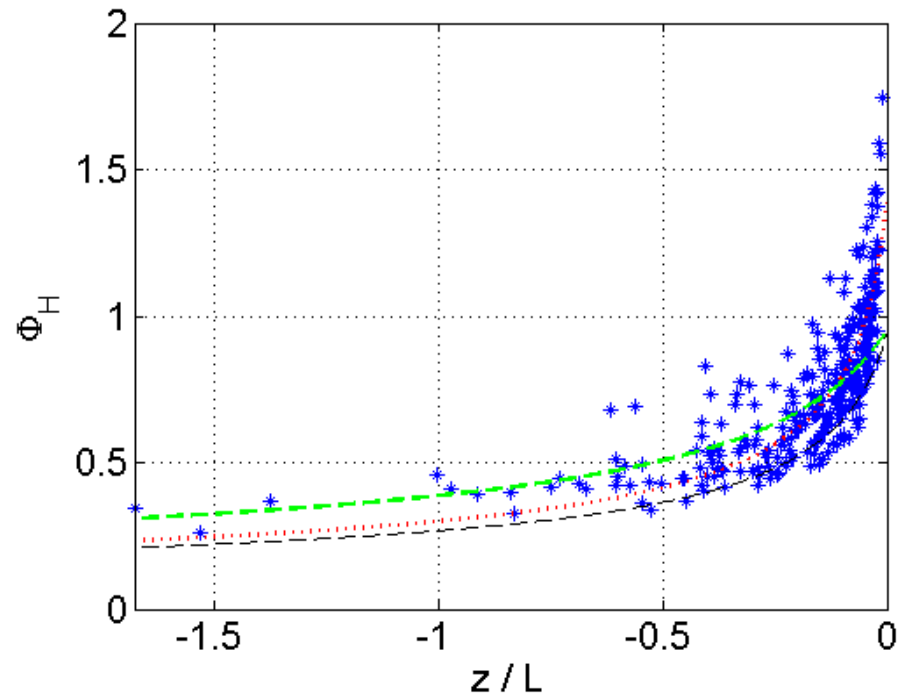
15 minutes



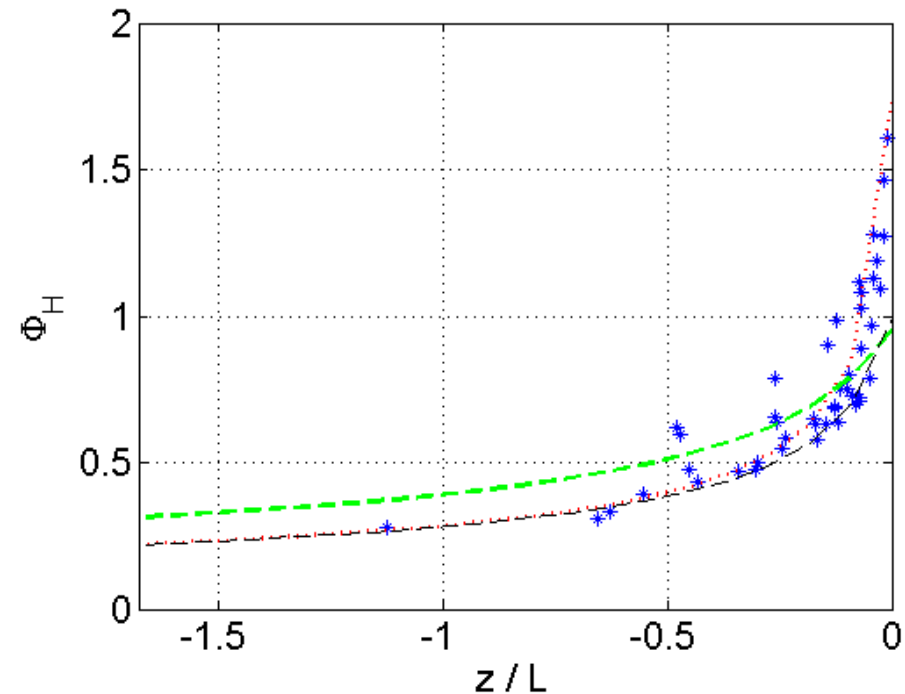
30 minutes

- Unstable case: Fitted curve adjusts to MOST for  $\zeta < -0.3$  but  $\Phi_H(0^-) \simeq 1.5$
- Stable case: The linear fit has a much smaller slope (22%), but  $\Phi_H(0^+) \simeq 0.95$

## Results (III): MOST for the 4 uppermost heights (Unstable)



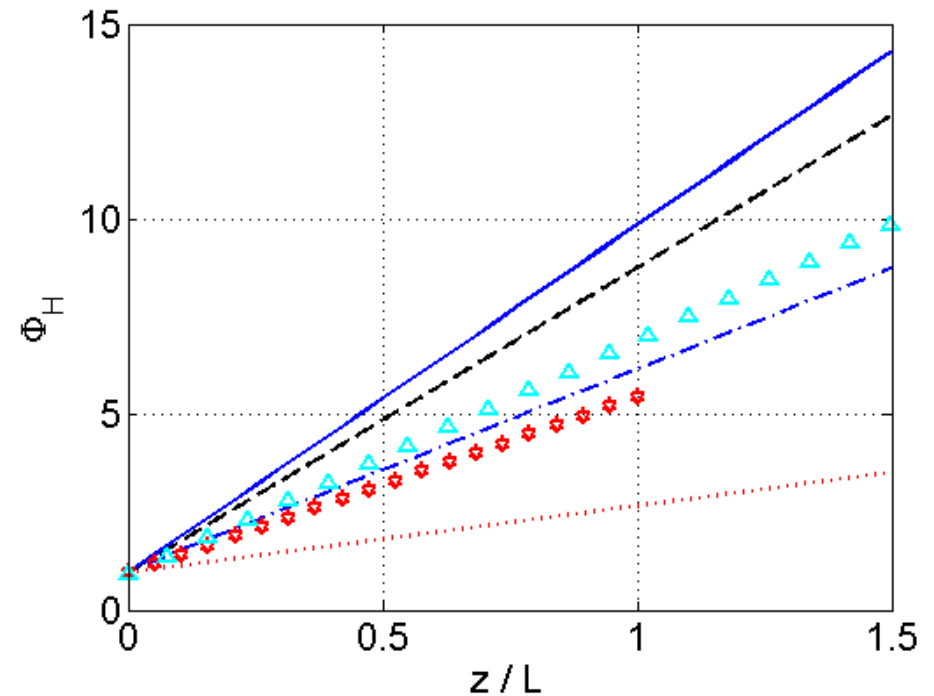
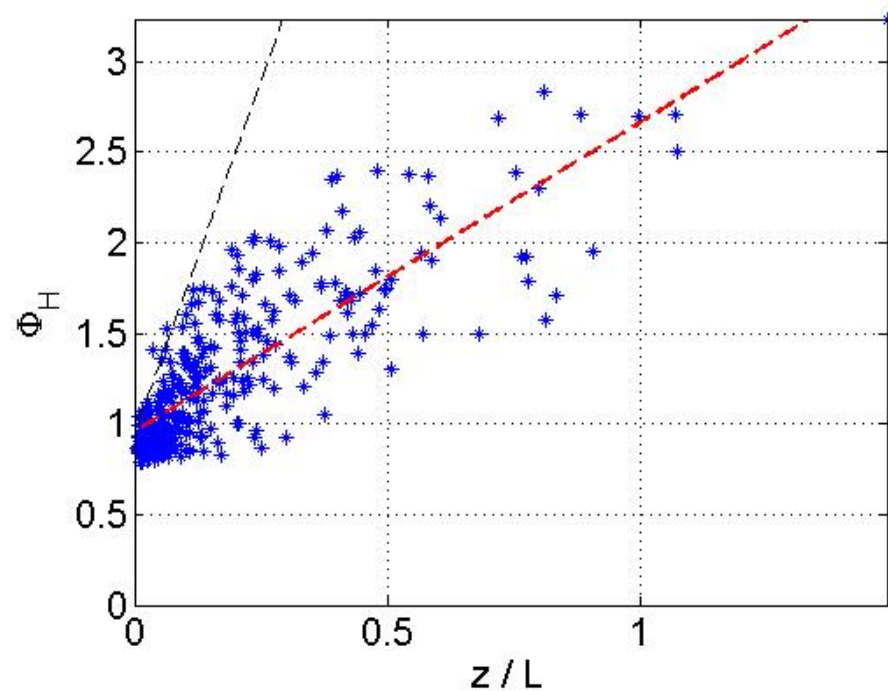
15 minutes



30 minutes

- A curved fitted to BLLAST data with  $\Phi_H(0^-) = 0.95$  do not provide better results.
- All the curves fitted with 15 min averaged data do not agree with the 30 min averaged data.

## Results (IV): MOST for the 4 uppermost heights (Stable)



Comparison of the results against different universal functions (from Foken, 2006).



## Sources of error

Högström (1988) summarized the possible reasons for the differences obtained experimentally:

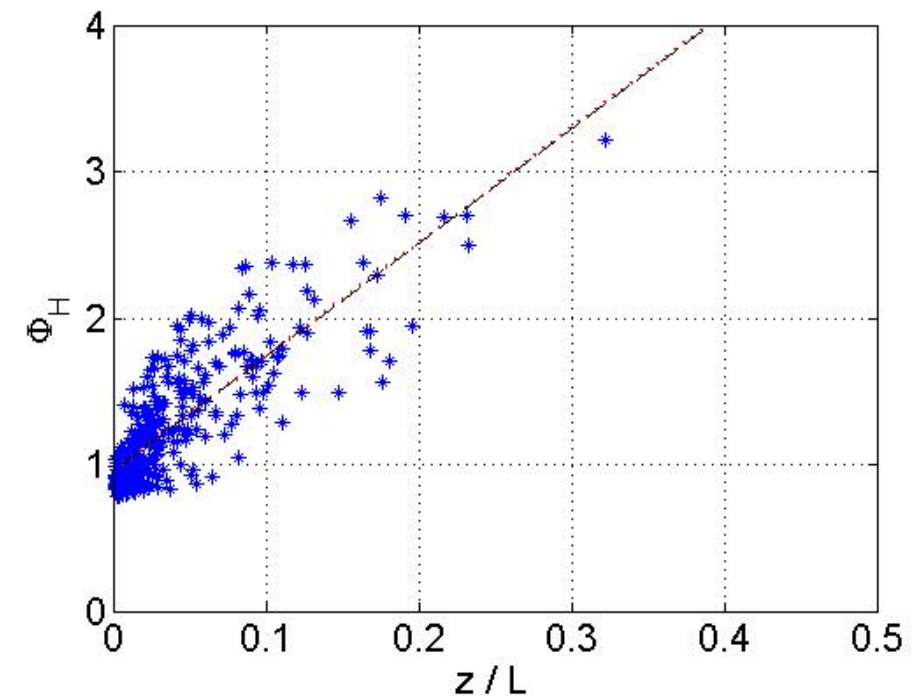
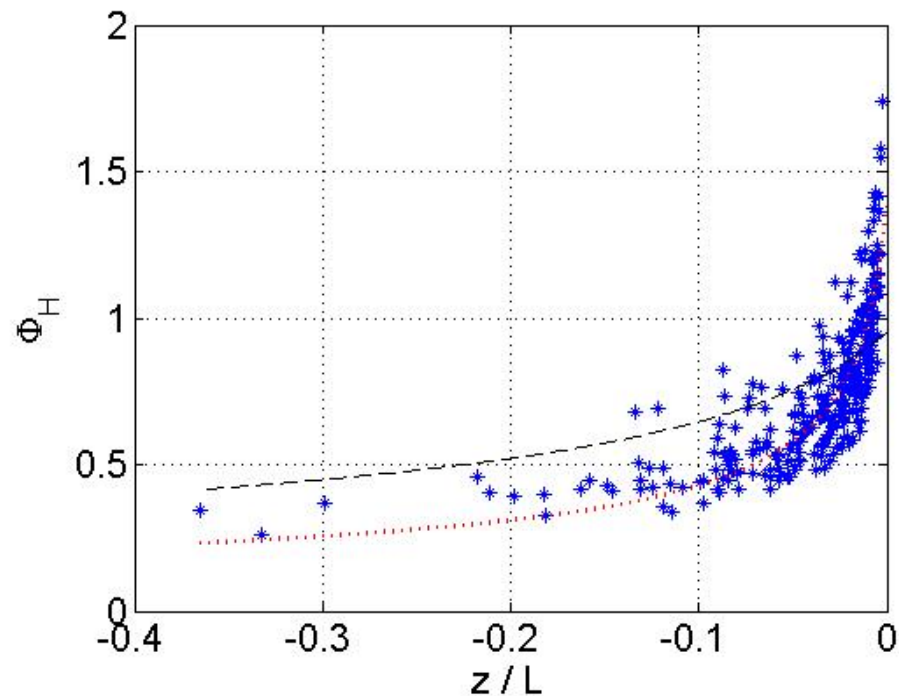
1. Statistical uncertainty in the individual averaging time periods
2. Systematic instrumental inaccuracy
3. Inadequate upwind fetch
4. Systematic errors due to inadequate sampling
5. Limitations of MOST

**Instrumental error:** Can we meet MOST with our dataset by correcting the flux measurements?

We define two independent correction factors ( $a$ ,  $b$ ):

$$\widetilde{\overline{w'\theta'}} = a \cdot \overline{w'\theta'} ; \quad \widetilde{u_*} = b \cdot u_*$$

## Results (V): MOST for the 4 uppermost heights (corrected)



- The averaged correction factors for the stable cases are  $a = 2.16$  and  $b = 2.15$ .
- The linear fit to the corrected values meets MOST for the stable cases.
- The results are not improved for the unstable cases.

## Preliminary conclusions

- The dataset provided by the 2-m column of thermocouples and the EC station at the SSH site do not follow MOST.
- Assuming horizontal homogeneity and steady conditions, these deviations can be produced by (i) systematic instrumental errors, (ii) uncertainty due to averaging time periods or (iii) limitations of MOST.
  - Unstable case: inadequate averaging period can not be discarded.
  - Stable case: A systematic error in the flux measurements could explain the deviation for heights  $z > 30$  cm, assuming that half of the flux is not recorded (due to smaller eddies?).
  - For the first 15 cm above ground, MOST cannot be followed. In this layer, other phenomena related with radiation flux divergence, latent heat processes, etc. must play a role.

## Acknowledgements

- BLLAST scientific and technical team for the support in the field.
- Joachim Reuder for the EC station data.