Evolution of the turbulence during the afternoon transition of the convective boundary layer



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Structure of the presentation



- 1. Summary of Darbieu et al. 2014
 - Research strategy
 - Data set
 - Methodology
 - Summary of the results

2. Confirmation/Invalidation of the assumptions

- Role of wind shear
- Impact of entrainement & demixing process
- Role of anisotropy
- Coherent structures

Motivation and research strategy



- > How do the turbulence characteristics vary in time and height?
- ➢ How does an idealized LES ?



Data used



BLLAST Campaign provides:

Turbulence observations along the afternoonConditions for an ideal case simulation

Large eddy simulation of an ideal case provides:

Complete vertical and temporal evolution of turbulence

An access to physical processes too complex to be observed

Vertical and temporal data coverage on 20 June 2011



+Surface energy balance stations

NCAR LES code (Moeng, 1984, 1986 / Sullivan et al., 1996 / Patton et al., 2005) •Case: 20 June 2011 of BLLAST campaign •Domain: 10 km x 10 km x 3 km Resolution: $\Delta x = \Delta y = 40$ m and $\Delta z = 12$ m •Prescribed initial profiles and surface flux based on most representative surface •Simple' wind profile •Prescribed total advection (AROME) •Similar thermal structure obtained •1.5 times less TKE in LES •Lower zi (850 m VS 1100 m)

Methodology: Spectral analytical model



Kristensen et al. 1989 : A generalized kinematic spectral model for anisotropic horizontally homogeneous turbulence $(1)^{2\mu}$



Results of Darbieu et al. 2014



LES ability to reproduce the turbulence evolution





The AT can be divided into two phases: the « Early Afternoon » and the « Late Afternoon »











Results of Darbieu et al. 2014





- > The changes occur first in the upper region of the PBL.
- > The higher within the PBL, the stronger the spectra shape changes.





















> New set of simulations





Dynamical production term











> TKE maintained in case of wind shear







> Wind shear delays the decay of TKE and maintains the TKE







- > Wind shear delays the decay of TKE and maintains the TKE
- Smaller TKE decay in the zones of mechanical production



- ➢ Sref: µ maintained until 1700 UTC
- ➤ U0: µ decreases everywhere from 1500 UTC
- ➢ U5: µ decreases earlier at the top of the PBL (no wind shear)





- > Sref: increase of I_w at the top of the PBL
- \succ U0: I_w constant
- \succ U5: I_w constant at the top of the PBL, decreases in the lower layers





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- Beta very similar (0.14)
- > U10 a bit larger in the morning
- U5 a bit smaller





The height at which the buoyancy flux is cancelling is very similar











- > Sref: According to I_w , the increase seems to propagate towards surface.
- The entrainement process might not be highlithed by studying buoyancy only.

H3: Could anisotropy modify the spectral characteristics?





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-0.2

-0.3

-0.4

-0.5

-0.6

-0.7

-0.8

-0.9

-0.2

-0.3

-0.4

-0.5

-0.6

-0.7

-0.8

-0.9

19

18



17 18 19

16

15

Time

0.3

12

13

14

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H4: Are the coherent structures at the origin of the steeper spectral slopes?





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Conclusions and perspectives





Conclusions and perspectives







Thank you for your attention























- ➢ New simulation initialized at 1100 UTC
- Same mean structure (Temperature, humidity, wind)
- Same zi
- Differences observed in the turbulent structure





 TKE starts to decrease everywhere at the same time





 TKE starts to decrease everywhere at the same time



The spectral changes happen earlier than in Darbieu et al. 2014





- ➢ New simulation starting at 1100 UTC
- Same mean structure (Temperature, humidity, wind)
- Same zi
- Differences observed in the **turbulent structure**:
 - > The TKE starts to decrease everywhere at the same time
 - > The spectral changes happen earlier than in Darbieu et al. 2014

LES does not have enough time to create the larger turbulent structures